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天文旋回约束下的页岩岩相分布与有机质富集 ——以四川盆地复兴地区东岳庙段为例

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摘要 【目的】四川盆地侏罗系陆相页岩油勘探取得了重大突破, 由于陆相泥页岩岩相粒度小、相变快和非均质性强等特点, 迫切需要对东岳庙段岩相发育规律及有机质富集特征开展系统研究。【方法】以四川盆地东部复兴地区自流井组东岳庙段为例, 结合岩心和非取心段岩相测井解释等资料, 基于旋回地层学理论, 建立了研究区高频等时地层格架, 通过天文旋回的沉积响应, 探讨了高频地层格架下泥页岩岩相的发育规律及有机质富集特征。【结果】(1)东岳庙段显示出良好的天文周期信号, 识别出长偏心率(405 ka)、短偏心率(128 ka)、斜率(43 ka)和岁差(21 ka)等天文旋回, 建立了4~5级等时地层格架。(2)研究区东岳庙段识别出2大类7亚类岩相, 两大类包括原地型富有机质纹层状泥页岩相、异地型含介壳(介壳质)泥岩相和介壳灰岩相。(3)东岳庙段主要受控于偏心率旋回, 405 ka和128 ka偏心率旋回驱动气候的变化控制了湖相含介壳泥页岩的沉积作用。【结论】长偏心率是控制岩相展布与有机质富集的关键因素, 主要影响了原地型富有机质纹层状泥页岩相、异地型含介壳泥岩相及介壳灰岩相的有序发育; 短偏心率对有机质富集影响有限, 但对异地型含介壳(介壳质)泥岩的岩相组合类型控制明显。

关键词 陆相页岩油; 岩相; 岩相组合; 天文旋回; 有机质富集; 四川盆地

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0 引言

近年来, 我国常规油气面临着接替储量不足, 产量持续下降, 发展困难等问题, “进(近)源找油”已成为必然趋势^[1]。页岩油气作为21世纪重要的油气接替领域, 是推动我国能源结构转型、降低油气对外依存度、缓解国内能源供需矛盾、保障国家能源安全的主力锚点与压舱石^[2-4]。2010年以来, 国内开始页岩油的勘探^[5], 目前已在准噶尔盆地吉木萨尔凹陷、鄂尔多斯盆地中部、江汉盆地潜江凹陷、渤海湾盆地济阳凹陷

和黄骅拗陷等地区实现了页岩油勘探的重大突破^[6-7]。

泥页岩的地层精细划分对比是约束岩相、沉积相分布与储层评价的重要基础。但是, 由于深水泥页岩普遍具有沉积粒度细、测井响应差异小的特点, 加之常规的层序地层学分析严重依赖垂向岩性组合变化或叠置样式, 导致深水泥页岩地层对比划分成为当前地层学研究最困难的领域之一^[8-11]。

基于米兰科维奇理论建立的旋回地层学, 以地球轨道周期为研究对象, 强调地球北纬65°夏季太阳日照量对地球气候的控制作用^[12-13]。尽管泥页岩地

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层中岩性变化小,但其内部矿物组成、有机质含量等参数具有极高的气候敏感性。借助地层剖面中测井、地球化学等连续数据的旋回地层学分析,为深水泥页岩的精细地层划分对比提供了新的思路^[7,14-17],这对于深水沉积区等时性高精度地层格架建立和岩相(及其组合)的分布预测、有机质富集规律的认识具有重要意义。

四川盆地东部复兴地区下侏罗统自流井组东岳庙段发育一套暗色富有机质页岩,具有分布稳定和有机质丰度高、类型好等特点,页岩油资源量大。前人针对其沉积环境、源岩特征、储集性能等方面进行研究^[18-22],明确了其形成于深湖一半深湖沉积背景,地层中粗粒物质主要来源于深水区周缘浅水介壳滩遭受的事件沉积作用,与深水原地沉积多样性混合发育了多种岩相及其组合类型。本文拟以四川盆地复兴地区东岳庙段为研究对象,通过旋回地层学分析识别长偏心率、短偏心率以及斜率等天文轨道周

期,指导井间多级次旋回地层对比,明确天文旋回约束下泥页岩岩相及其组合发育和有机质富集的规律,旨在提高页岩油优质储层(甜点)预测精度,改进水平井轨迹设计方案。

1 地质背景

复兴地区位于四川盆地东部,主体处于重庆梁平、垫江和丰都,面积约3 000 km²,属于北东向川东高陡断褶带南部地区,平面上呈“两隆三洼”的构造格局,东西被大池干背斜和明月峡背斜带所夹持,中部被黄泥塘和苟家场背斜分割为梁平向斜、拔山寺向斜和大堡场向斜^[23](图1a)。自加里东期至印支早期,四川盆地主要以海相的碳酸盐岩沉积为主,晚三叠世逐渐由海相沉积转为陆相沉积,早侏罗世四川盆地进入构造活动的宁静期和短暂的陆内伸展阶段。四川盆地在下侏罗统自流井组依次发育珍珠冲段、东岳庙段、马鞍山段和大安寨段(图1b)。珍珠冲

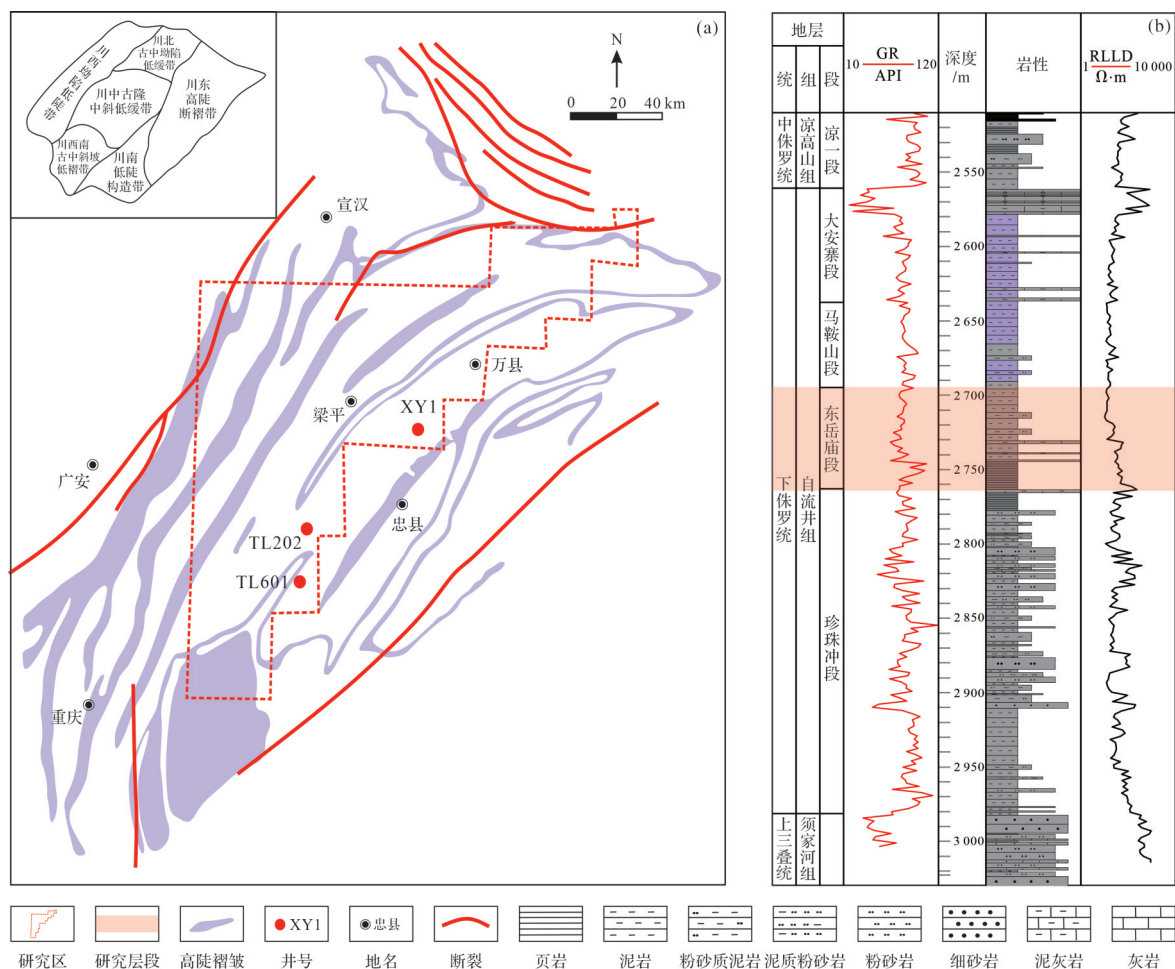


图1 复兴地区位置及XY1井自流井组地层柱状图(据文献[24]修改)

Fig.1 Location of the Fuxing area and stratigraphic column of the Ziliujing Formation in well XY1 (modified from reference [24])

段处于湖盆扩张早期,快速沉降背景下发育了规模较大的三角洲—滨浅湖沉积体系;东岳庙段沉积期盆地基底沉降速度变缓,周围构造带活动逐渐停息,物源供应缓慢,湖平面的上升,加剧了大规模的深水湖相沉积的发育;马鞍山段沉积时期伸展作用变强,盆地快速、稳定沉降,物源供应充足,盆缘发育规模较大的河流—三角洲沉积;大安寨段盆缘造山等构造活动再次平息,盆地稳定沉降,湖盆扩张到最大^[23]。

2 数据与方法

研究使用的数据来自XY1井、TL202井和TL601井东岳庙段的自然伽马(GR)、自然伽马能谱(NGR)、矿物元素(ECS)等测井数据,XY1井和FY10井的稳定碳同位素数据以及其取心段完成的主量元素、总有机碳(TOC)含量和全岩矿物测试数据。其中,稳定碳同位素数据采用DELTA PLUS V稳定同位素质谱仪(YQ3-12-13 编号)依据国标 GB/T 18340.2—2010进行测试,主量元素采用 AxiosmAX X 射线荧光仪依据国标 GB/T 14506.28—2010 完成测试,TOC 含量采用 CS844 碳硫分析仪器(3250 编号)完成测试,X 射线衍射全岩分析采用 D/max-2600 型 X 射线衍射仪(9005)仪器,依据 SY/T5163-2018《沉积岩中黏土矿物和常见非黏土矿物 X 射线衍射分析方法》进行测定。选取复兴地区位于湖盆深水区的有丰富取心资料的井和周边无取心资料的井作为研究对象开展研究。

作为古气候替代指标之一,GR 高值指示黏土和有机质含量较高,气候湿润,化学风化和陆源输入更强^[25]。本文利用选取 XY1 井、TL202 井和 TL601 井采样间距为 0.125 m 的 GR 数据采用软件 Acycle v2.1 版本^[26]开展旋回地层学研究。通过对天文轨道理论模型进行分析,获取目的层段地质时期的天文轨道参数;其次进行时间序列分析,将数据序列由深度域转化为频率域,根据频率域的能量强度与置信水平识别由天文轨道参数控制的频率信号^[25-26]。具体过程包括数据预处理、频谱分析、滤波分析以及天文调谐等。本次研究使用预白化处理方式,采纳“LOESS”、“LOWESS”等局部回归平滑方法去除长趋势,增强高频信号压制低频信号^[26];通过 COCO 相关系数法即零假设的显著性检验和蒙特洛迭代确定最佳沉积速率,减少了人为因素^[27];使用“MTM”多窗谱分析方法将数据序列由深度域转换为频率域,进行频谱分析获取在不同级别天文轨道参数控制下旋回地层厚度

的置信区间^[28];通过滑动频谱分析,检验沉积地层中可能存在的沉积间断和沉积速率的变化;通过带通滤波提取出深度域数据序列中的波形曲线,进而提取偏心率信号、斜率信号和岁差信号。

此外,利用 ECS 数据(间距 0.15 m)辅助 XY1 井岩心资料开展岩相及其组合判别。最后,在取心段实测基础上,为了克服实测样品采样的不均匀,总有机碳的含量数据采用实测数据约束下 $\Delta\lg R$ 改进的测井解释模型^[29]求取。

3 研究结果

3.1 岩相类型及沉积特征

基于 XY1 井的岩石薄片、X 射线衍射全岩分析及测录井资料,采用结构优先原则,以组分和沉积构造为主要依据,将研究区东岳庙段细粒沉积岩识别出 2 大类 7 亚类岩相(图 2)。其中,两大类包括:(1)原地型富有机质纹层状泥页岩相、(2)异地型含介壳(介壳质)泥岩相和介壳灰岩相;7 亚类包括:(1)纹层状富粉砂黏土质泥岩相、(2)块状富粉砂黏土质泥岩相、(3)纹层状混合质泥岩相、(4)纹层状含介壳富粉砂黏土质泥岩相、(5)纹层状含粉砂富介壳黏土质泥岩相、(6)纹层状含粉砂富黏土介壳质泥岩相和(7)块状泥质介壳灰岩相(表 1)。

1) 原地型富有机质纹层状泥页岩相

纹层状富粉砂黏土质泥岩相岩心上以灰黑色为主,贫介壳,纹层较发育,镜下可以观察到浅色的粉砂质纹层和深色的黏土质纹层相间,成层性较好(图 3a、图 4a);块状富粉砂黏土质泥岩相岩心观察为灰黑色,均质块状,贫介壳,镜下黏土含量较高,颗粒杂乱堆积无定向,层理不发育(图 3b、图 4b);纹层状混合质泥岩相岩心上为深灰色,均质块状,贫介壳,镜下呈点一线接触,半定向分布,成层性较差(图 3c、图 4c)。

2) 异地型含介壳(介壳质)泥岩相及介壳灰岩相

纹层状含介壳富粉砂黏土质泥岩相岩心上呈灰色,可见少量介壳纹层,介壳保存完整,镜下呈线接触,介壳层定向分布,成层性较好(图 3d、图 4d);纹层状含粉砂富介壳黏土质泥岩相岩心上呈灰色,介壳纹层发育且保存完整,镜下呈线接触,可见原生和经成岩改造的两种介壳(图 3e、图 4e);纹层状含粉砂富黏土介壳质泥岩相岩心上为深灰色,无介壳,镜下见泥晶方解石屑,与富有机质黏土质叠合(图 3f、图 4f);

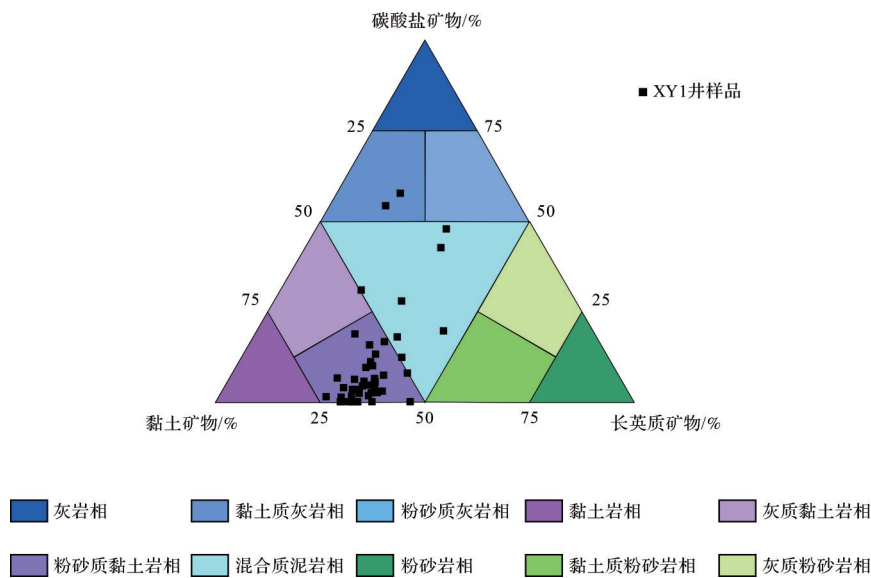


图2 复兴地区东岳庙段矿物组分三端元

Fig.2 Three terminal elements of mineral components in the Dongyuemiao member in the Fuxing area

表1 复兴地区东岳庙段岩相类型及其沉积特征

Table 1 Lithofacies and sedimentary characteristics in the Dongyuemiao member, Fuxing area

岩相亚类	长英质矿物/%	碳酸盐矿物/%	黏土矿物/%	沉积构造*
纹层状富粉砂黏土质泥岩相	25~50	<25	>50	纹层状
块状富粉砂黏土质泥岩相	25~50	<25	>50	块状
纹层状混合质泥岩相	25~50	25~50	25~50	纹层状
纹层状含介壳富粉砂黏土质泥岩相	25~50	10~25	>50	纹层状
纹层状含粉砂富介壳黏土质泥岩相	10~25	25~50	>50	纹层状
纹层状含粉砂富黏土介壳质泥岩相	10~25	>50	25~50	纹层状
块状泥质介壳灰岩相	—	>50	—	块状

注:纹层状厚度小于1 cm;厚层状厚度介于1~10 cm;块状厚度大于10 cm。

块状泥质介壳灰岩相岩心上为灰白色,富介壳,破碎接触,均质块状,镜下可见石英充填部分介壳间的粒间孔(图3g,h)。

3.2 旋回地层划分

3.2.1 地质年龄与理论天文周期

前人在英国、法国、意大利、葡萄牙以及中国等地区均发现S-P事件,认为S-P事件具有全球性,约193 Ma^[30,31]。通过对比稳定碳同位素曲线,发现四川盆地东北部陆相东岳庙段 $\delta^{13}C_{PDB}$ 曲线与同期海相记录具有显著的相似性。在辛涅缪尔阶和普林斯巴阶边界事件(S-P事件)的碳同位素负偏移幅度是变化的,英国约3‰,法国约6‰,四川盆地约8‰^[30],英国Cardigan Bay盆地和法国Paris盆地在辛涅缪尔阶和普林斯巴阶有较为连续的、年龄约束较好的碳同位素记录^[32-33](图5)。

通过Acycle软件获取了La2004天文轨道方案中180~201 Ma之前的ETP数据,并使用MTM方法对ETP数据进行频谱分析,分别是长偏心率E,短偏心率e1和e2,斜率O1和O2,岁差P1、P2和P3。据此偏心率、斜率、岁差周期的比值表,可知理论天文轨道参数的周期近似满足23.1:7.3:5.6:2.5:2.0:1.2:1.1:1.0。

3.2.2 天文旋回识别

本次研究采用“LOWESS”局部回归平滑方法进行去趋势处理,对XY1井的GR数据原始序列去除35%的长期趋势,使数据序列均值为0。

在鲁棒红躁模型Robust AR(1)下,采用MTM方法,对去趋势后的GR数据进行频谱分析。频谱图中90%置信度曲线之上,存在多个能量较高的频率峰值,代表的旋回厚度约13.8 m、10.8 m、8.4 m、4.8 m、

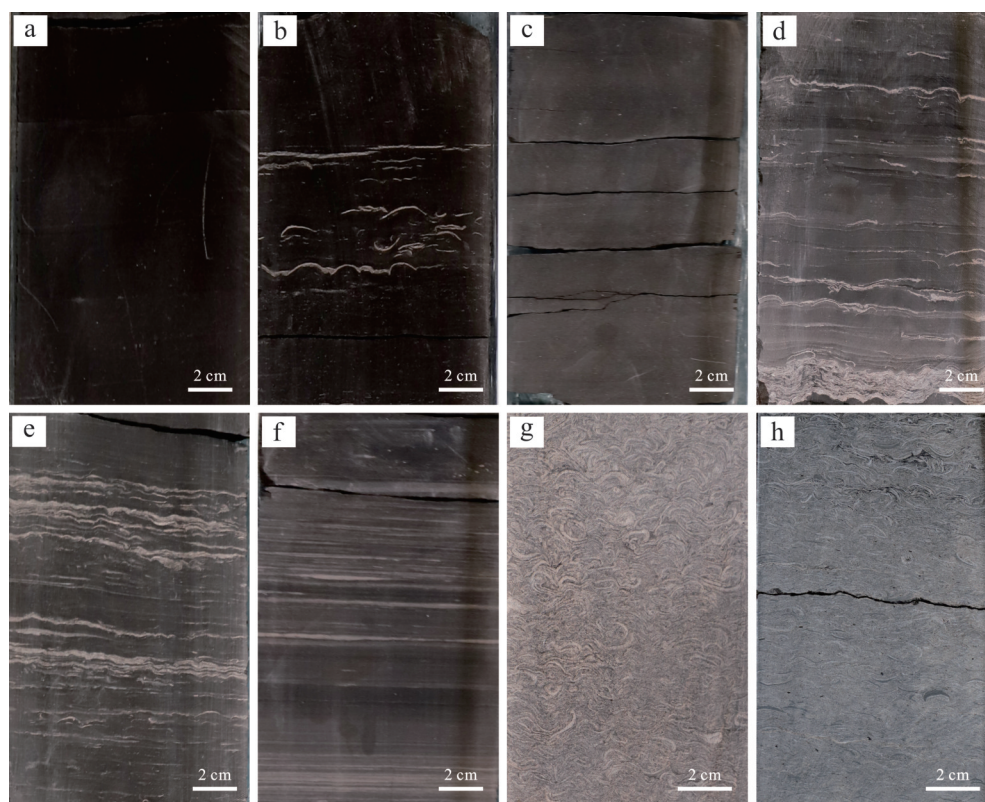


图3 复兴地区东岳庙段岩心中不同岩相沉积特征(XY1井)

(a)纹层状富粉砂黏土质泥岩相, 2 864.4 m; (b)块状富粉砂黏土质泥岩相, 2 863.6 m; (c)纹层状混合质泥岩相, 2 831.5 m; (d)纹层状含介壳富粉砂黏土质泥岩相, 2 840.1 m; (e)纹层状含粉砂富介壳黏土质泥岩相, 2 856.1 m; (f)纹层状含粉砂富黏土介壳质泥岩相, 2 836.4 m; (g)块状泥质介壳灰岩相, 2 843.0 m; (h)块状泥质介壳灰岩相, 2 821.6 m

Fig.3 Core characteristics of different lithofacies in the Fuxing area (well XY1)

(a) laminated silty-rich argillaceous mudstone facies, 2 864.4 m; (b) massive silt-rich argillaceous mudstone facies, 2 863.6 m; (c) laminated mixed mudstone facies, 2 831.5 m; (d) laminated shell-bearing silty-rich argillaceous mudstone facies, 2 840.1 m; (e) laminated silt-bearing shell-rich argillaceous mudstone facies, 2 856.1 m; (f) laminated silt-bearing clay-rich shelly mudstone facies, 2 836.4 m; (g) massive argillaceous shell limestone facies, 2 843.0 m; (h) massive argillaceous shell limestone facies, 2 821.6 m

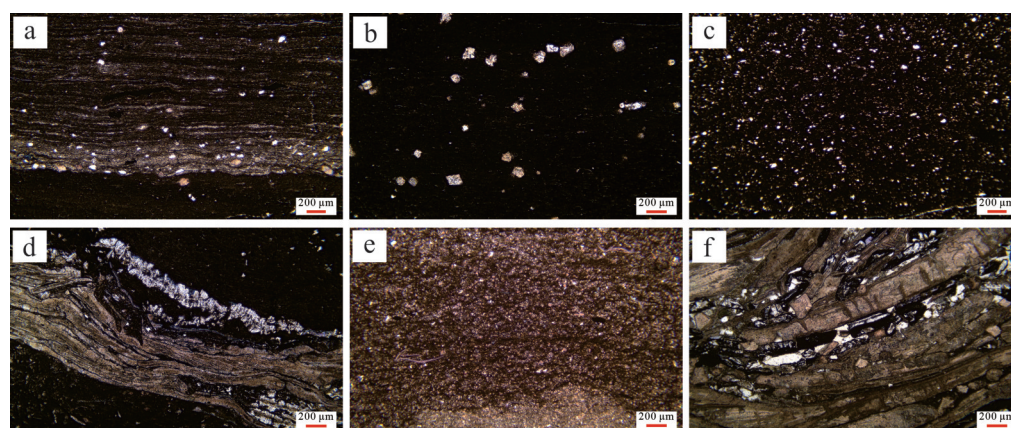


图4 复兴地区东岳庙段不同岩相镜下薄片特征(XY1井)

(a)纹层状富粉砂黏土质泥岩相, 2 864.4 m; (b)块状富粉砂黏土质泥岩相, 2 863.6 m; (c)纹层状混合质泥岩相, 2 831.5 m; (d)纹层状含介壳富粉砂黏土质泥岩相, 2 840.1 m; (e)纹层状含粉砂富介壳黏土质泥岩相, 2 856.1 m; (f)纹层状含粉砂富黏土介壳质泥岩相, 2 836.4 m

Fig.4 Thin slice characteristics under different lithofacies in the Fuxing area (well XY1)

(a) laminated silty-rich argillaceous mudstone facies, 2 864.4 m; (b) massive silt-rich argillaceous mudstone facies, 2 863.6 m; (c) laminated mixed mudstone facies, 2 831.5 m; (d) laminated shell-bearing silty-rich argillaceous mudstone facies, 2 840.1 m; (e) laminated silt-bearing shell-rich argillaceous mudstone facies, 2 856.1 m; (f) laminated silt-bearing clay-rich shelly mudstone facies, 2 836.4 m

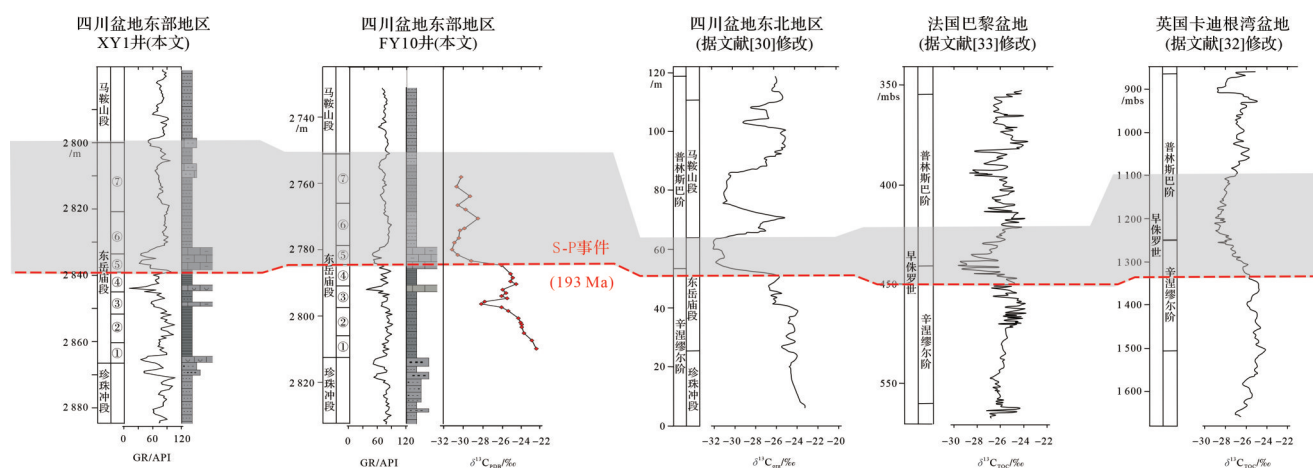


图5 四川盆地早侏罗世东岳庙段与其他盆地碳同位素垂向变化对比与S-P事件(193 Ma)的确定

Fig.5 Comparison of vertical changes of carbon isotopes between the Dongyuemiao member and other basins in the Early Jurassic and determination of the S-P event (193 Ma) in the Sichuan Basin

4.0 m、3.3 m、2.2 m和1.0 m,其中13.8:8.4:4.8:2.2与天文周期理论比值接近,认为该套沉积地层记录了405 ka长偏心率E、128 ka短偏心率e1、98 ka短偏心率e2和43 ka斜率O1等多级别天文信号(图6a)。EHA滑动能谱图显示,405 ka长偏心率在深度域中连续性较好,较为稳定,但在2 830 m处存在左偏,表明沉积速率增大(图6b),相应的旋回控制的沉积地层厚度增大,其他的天文信号连续性较差,断续显示,由于整套沉积地层厚度较大,因此对天文信号的显示与提取存在一定的误差。

通过相关系数分析(COCO)方法对去趋势后的GR数据进行零假设显著性检验,确定XY1井的最佳平均沉积速率。采用Pearson方法,实验沉积速率范围设置为1~5 cm/ka,实验间隔设置为0.1 cm/ka,蒙特卡洛迭代次数设置为5 000次,尼奎斯特频率(Nyquist frequency)设置为5,去除鲁棒红噪模型。结果显示,零假设显著性主要存在2.7 cm/ka、3.5 cm/ka、4.3 cm/ka三个峰值,说明地层可能存在三个主要沉积速率,而相关系数存在3.5 cm/ka一个主峰,综合二者可知3.5 cm/ka的零假设显著性最低,为最佳平均沉积速率(图6c)。通过带通滤波提取出深度域数据序列中的13.8 m、4.8 m、2.2 m、1.0 m等对应的天文信号曲线,由于405 ka长偏心率周期在地质历史时期相对稳定,可以利用405 ka的周期完成调谐过程,使得天文信号由深度域转为时间域,将旋回提取的旋回周期与天文轨道周期相对比,推测东岳庙段沉积持续时间约1.63 Ma,在自流井组东岳庙段XY1井主

要识别出4个405 ka长偏心率旋回,15个128 ka短偏心率旋回(图6d)。依据上述方法,对TL202井和TL601井进行单井旋回地层学分析,分别提取出了4个405 ka长偏心率周期和15个128 ka短偏心率周期。

3.2.3 天文旋回划分结果

目前大多数学者认为,高频层序受到地球轨道周期性变化的天文因素所控制,其中四级层序和五级层序分别与长偏心率周期和短偏心率周期对应^[34]。依据GR曲线对环境的指示,通常将天文周期的极小值作为基准面的界限,即高频层序的界面。本文将天文周期的极小值为主要划分依据,结合测井曲线、岩心岩性和总有机碳含量等数据,进行四级层序和五级层序的划分。在高精度地层单元识别与划分基础之上,遵循“分级对比,逐级控制”的基本原则,四级层序约束五级层序,利用最大洪泛面为标志层,分析高频层序在空间上的变化,逐一进行对比。据此,将XY1井、TL202井和TL601井东岳庙段均划分出4个四级层序(E1~E4)和15个五级层序(e1~e15)。结果显示,每口井的E1、E3和E4长偏心率旋回均包含4个e短偏心率旋回。

4 天文旋回约束下的岩相展布与有机质富集特征

4.1 古气候、矿物组成对天文旋回的响应

鉴于东岳庙段1~5小层为页岩油勘探开发的重

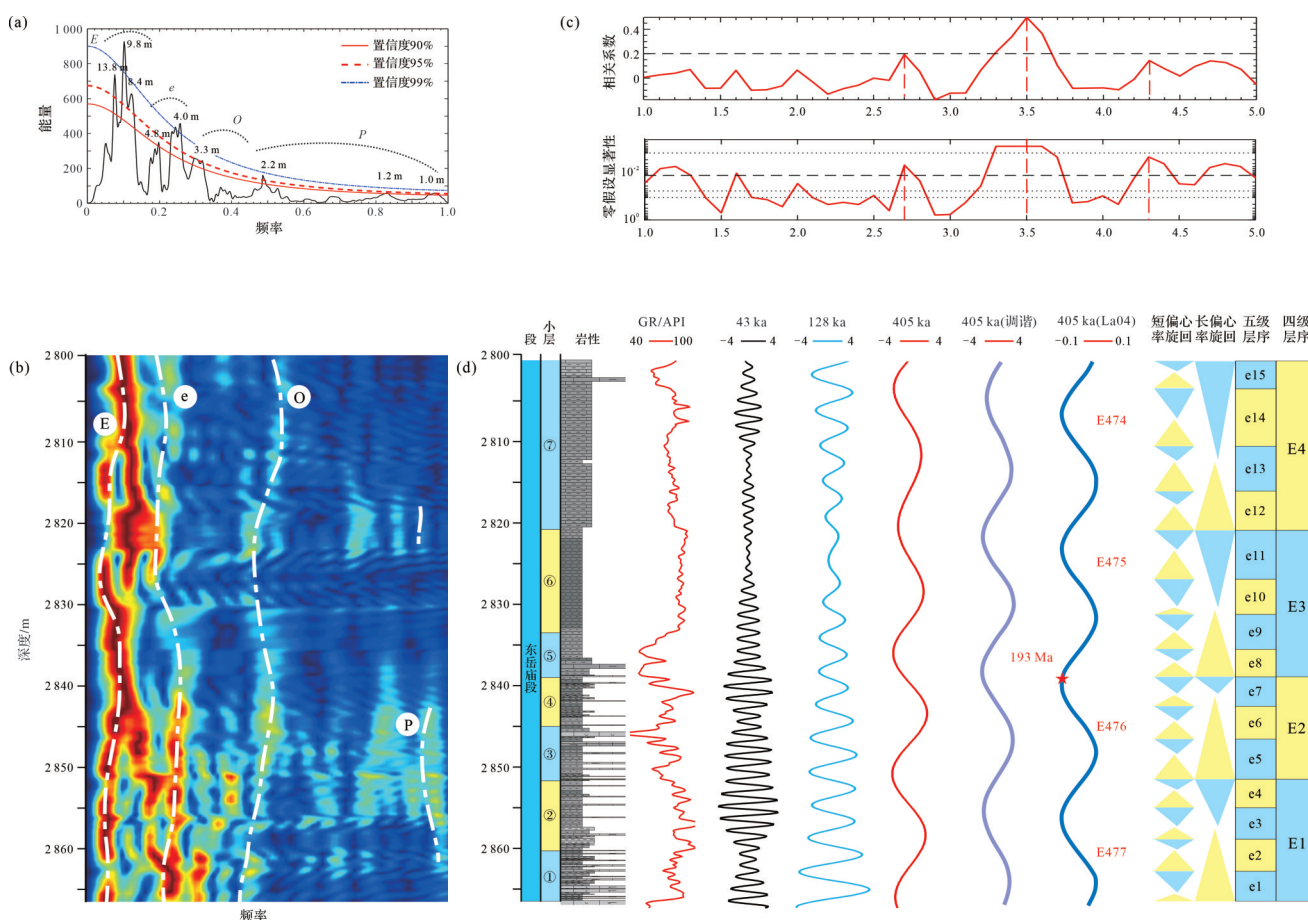


图6 XY1井东岳庙段旋回地层分析结果

(a)GR数据的频谱分析结果;(b)GR数据的滑动频谱分析结果;(c)GR数据的相关系数分析结果(COCO);(d)GR数据的各级周期滤波提取以及层序划分结果

Fig.6 Cyclic stratigraphic results of the Dongyemiao member in well XY1

(a) spectral analysis results of gamma ray (GR) data; (b) sliding spectrum analysis results of GR data; (c) correlation coefficient analysis results of GR data (COCO); (d) the GR data were extracted by periodic filtering at all levels and the results of sequence division

点层段,因此选取其开展页岩岩相展布及有机质富集研究。通过古气候干湿指数(C值)、Ni/Co、Sr/Ba、Fe/Mn值等元素指标来判别沉积环境。研究认为古气候C值大于0.8、介于0.6~0.8、0.4~0.6、0.2~0.4、小于0.2分别表明古气候为湿润、半湿润、半干旱到半湿润、半干旱和干旱环境^[35]。东岳庙段C值介于0.07~2.32,平均值为0.95,反映出沉积期的古气候以湿润为主;Ni/Co值小于5指示富氧环境,Ni/Co介于5~7指示贫氧环境,Ni/Co值大于7指示准厌氧—厌氧环境^[36]。Ni/Co值介于0.44~1.19,平均值为0.69,整体为富氧环境;Sr/Ba值小于0.5指示淡水环境,Sr/Ba值介于0.5~1.0指示半咸水环境,Sr/Ba值大于1指示咸水环境^[37]。Sr/Ba值介于0.20~1.68,平均值为0.37,其中东岳庙段1~4小层Sr/Ba平均值为0.26,5小层Sr/Ba平均值为0.74,反映出东岳庙段1~4小层沉积于淡水环境,进入5小层后水体过渡为咸水;Fe/Mn值

小于30为深湖沉积,介于30~50为半深湖沉积,大于50为浅湖沉积^[38]。Fe/Mn值介于33.72~219.03,平均值为146.98,整体处于半深湖沉积。整体上而言,东岳庙段1~5小层气候由温暖湿润、富氧、淡水的半深湖环境,进入5小层后,水体开始咸化。将长、短偏心率曲线与各环境指标相拟合,发现短偏心率曲线与古水深、氧化还原条件变化趋势确实相对一致。即短偏心率极大值附近,湖平面高,水体还原性更强,而短偏心率极小值附近,湖平面低,水体还原性较弱。因此,认为研究区东岳庙段的气候以及沉积环境受到偏心率的驱动。

陆源碎屑矿物(石英+长石)含量可以反映风化作用的强度,ECS矿物测井显示,XY1井长英质含量最大为49.5%,最小为9.6%,平均值为35.2%,表明复兴地区在东岳庙段陆源输入较强(图7)。其中,长偏心率极大值附近的黏土矿物与长英质矿物含量较

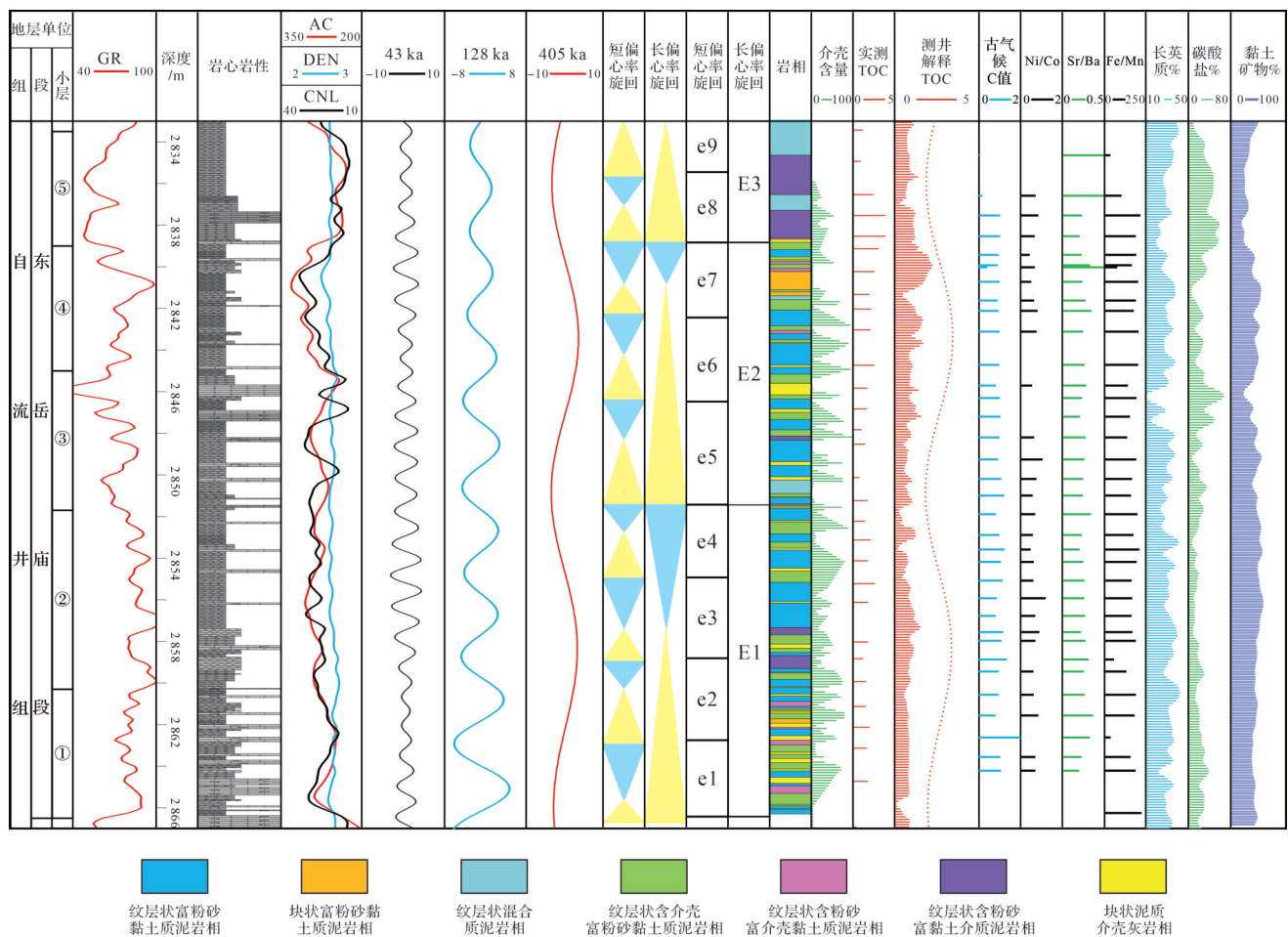


图7 XY1井东岳庙段1~5小层长偏心率、短偏心率旋回划分及其约束下的岩相特征及有机质富集特征

Fig.7 Division of the long and short eccentricity cycles and their lithofacies and organic matter enrichment characteristics under the constraint of layers 1-5 in the Dongyue Temple member of well XY1

多,而碳酸盐矿物的含量较少。相反,在长偏心率极小值附近,碳酸盐矿物及介壳的含量增多,而黏土矿物与长英质矿物的含量减少。除长偏心率周期与矿物组分含量具有较好的对应关系外,短偏心率周期在更小尺度上也具有同样的地质响应特征。因此,偏心率越大,偏心率振幅越大,气候岁差振幅越大,干湿变化越明显。显著的气候即干湿变化会导致更强的物理和化学风化作用,在强季风雨降条件下,水体变深,会产生大量的细粒物质和长英质等陆源碎屑,并携带大量的营养物质^[38-39],因此,藻类的光合作用变强,碳酸盐矿物沉淀增强。

4.2 长偏心率旋回约束下的岩相与有机质分布

东岳庙段E1、E2两个四级层序整体发育半深湖—深湖相。东岳庙段各岩相在XY1井(图7)四级层序E1和E2均有分布,其中E1下部以异地型含介(介壳质)泥岩及介壳灰岩相为主,多种岩相类型交

替出现。最大洪泛面附近原地型富有机质纹层状泥页岩较为发育,随着进一步湖退,异地型含介(介壳质)泥岩及介壳灰岩相逐渐增多;E2整体以原地型富有机质纹层状泥页岩相为主,其中异地型含介(介壳质)泥岩及介壳灰岩相夹层自下而上逐渐减薄;E3开始大范围湖退,底部发育厚层的异地型介壳质泥岩。

在XY1井—TL202井—TL601井北东—南西向剖面上(图8),E1、E2在XY1井附近水体较浅,发育大量生物介壳,沉积了含介壳(介壳质)泥岩及介壳灰岩相,逐渐向湖盆边缘减薄。最大洪泛面附近生物介壳罕见,均以原地型富有机质纹层状泥页岩相为主,厚度较为稳定;古物源和古水深一定程度上影响陆源碎屑的输入和生物碳酸盐的产量,进而控制岩相的类型,古物源与古地貌控制了岩相的分布。通常,在水体较浅清澈、多物源、靠近湖盆边缘的环

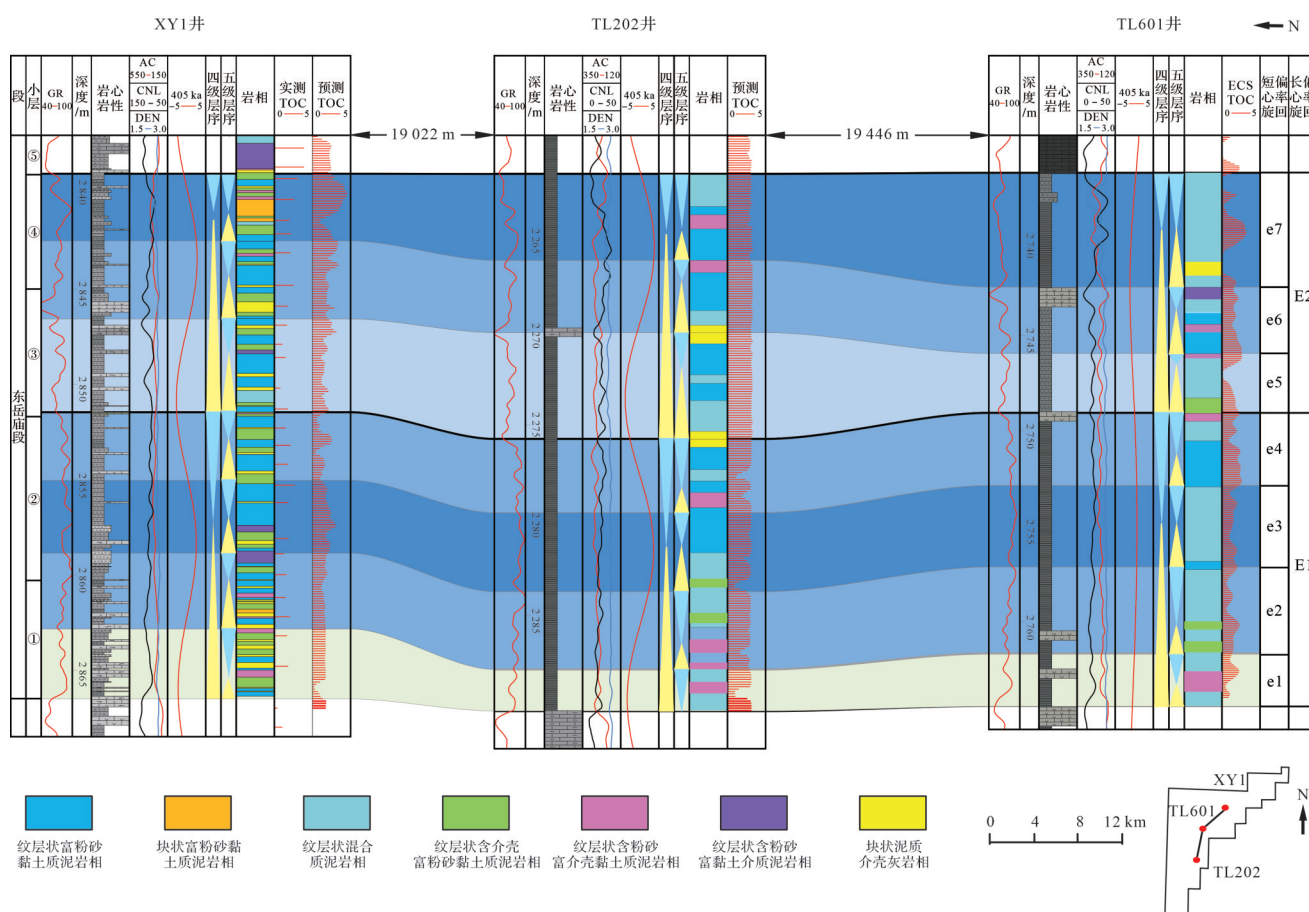


图8 复兴地区长偏心率约束下东岳庙段1~5小层对比及岩相展布特征

Fig.8 Stratigraphic correlation and lithofacies distribution characteristics of layers 1-5 in the Dongyuemiaomiao member constrained by long eccentricity in the Fuxing area

境下,生物介壳繁盛,碳酸盐含量较高,异地型含介(介壳质)泥岩及介壳灰岩相比较发育;而在水体较深、少物源、还原性较强、靠近湖盆中心的环境下,黏土含量较高,粉砂质页岩与黏土质页岩相对发育,以原地型富有机质纹层状泥页岩相为主。通常认为,天文旋回驱动古气候影响研究区沉积地层的矿物含量,进而控制岩相类型^[13,40]。由此可见,原地型和异地型泥岩相的展布严格受到四级层序的控制,在长偏心率极大值附近发育原地型富有机质纹层状泥页岩相,相反在长偏心率极小值附近发育异地型含介(介壳质)泥岩及介壳灰岩相。

东岳庙段e6水体较深,发育半深湖—深湖相。在XY1井(图8)附近,e6底部生物介壳繁盛,发育大套厚层的异地型含介(介壳质)泥岩及介壳灰岩相,随着水体进一步加深,生物介壳含量变少,异地型岩相夹层厚度逐渐变薄。在XY1井—TL202井—TL601井北东—南西向剖面上(图8),向盆地边缘方

向,生物介壳变多,异地型岩相夹层厚度逐渐变厚。由此可见,异地型泥岩及介壳灰岩相的展布严格受到五级层序的控制,在短偏心率极小值附近发育异地型含介(介壳质)泥岩及介壳灰岩相,随着短偏心率增大,异地型含介(介壳质)泥岩及介壳灰岩相厚度逐渐减薄。结合TOC曲线与天文旋回信号,发现短偏心率在小尺度范围内与TOC的相关性较弱,对有机质富集控制有限。

研究区有机质类型以Ⅱ₂型为主^[20-21],其来源主要为陆生的高等植物,通过陆源输入进入湖盆沉积,长英质矿物、黏土矿物也反映物源输入。因此,TOC高值指示长英质矿物、黏土矿物和有机质含量高,反映水体较深的沉积环境,对应纹层状构造,反之TOC低值指示长英质矿物、黏土矿物和有机质含量低,反映水体较浅的沉积环境,对应弱纹层状构造和块状构造。此次研究发现,TOC曲线与沉积环境指标变化一致,二者密切相关,说明有机质富集受控于沉积

环境的变化^[41]。通常,温暖湿润的气候降雨充沛,纹层状和层状构造较为发育,河流可以携带大量的陆源碎屑进入湖盆,并且水生生物和藻类生物繁盛,营养物质增多,湖泊的古生产力增强,使得有机质更加富集^[42-44]。偏心率是地球绕太阳公转轨道面偏离正圆的程度,也可以通过调制岁差周期的变幅使得季风的周期性变化^[44]。天文旋回控制了古气候,进而影响研究区湖相泥页岩的沉积^[45]。前人研究发现,东营凹陷沙三下亚段总有机碳受到偏心率和斜率的控制^[41]。结合405 ka长偏心率和128 ka短偏心率天文周期曲线,发现总有机碳含量与405 ka长偏心率周期趋势一致,长偏心率极大值附近TOC值越大,推测地球轨道参数长偏心率驱动气候的变化,进而控制了有机质的富集(图9)。

5 结论

(1) 通过米兰科维奇旋回对XY1井、TL202井、

TL601井进行高频旋回识别,分别识别出4个405 ka的长偏心率旋回(四级层序)、15个128 ka短偏心率旋回(五级层序),建立了复兴地区高频等时地层格架。

(2) 基于矿物组成和沉积构造,建立了复兴地区东岳庙段湖相深水泥页岩岩相划分方案,共识别出原地型富有机质纹层状泥页岩相和异地型含介壳(介壳质)泥岩相和介壳灰岩相两大类岩相。

(3) 原地型富有机质纹层状泥页岩相多发育在温暖湿润、水深较大、偏还原环境的淡水安静环境,对应于偏心率极大值附近,而异地型含介壳泥岩相常存在于干旱寒冷、水深较浅、偏氧化的盐度较大的动荡环境,对应于偏心率极小值附近。鉴于天文旋回对岩相与有机质的响应,综合分析发现长偏心率是控制岩相展布与有机质富集的关键因素,主要影响了原地型富有机质纹层状泥页岩相和异地型含介壳泥岩相的发育;短偏心率对有机质富集影响有限,

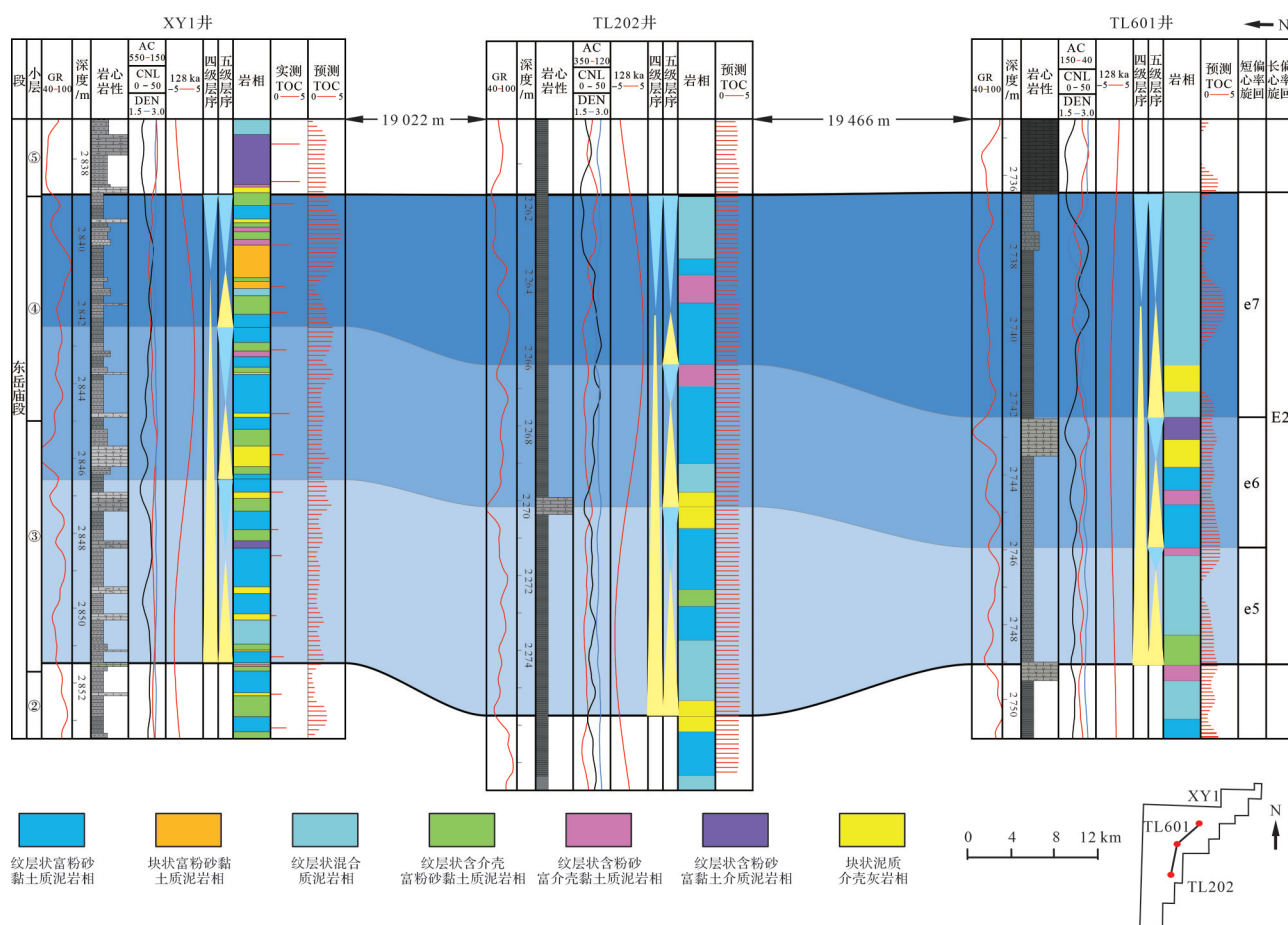


图9 复兴地区短偏心率约束下东岳庙段1~5小层对比及岩相展布特征

Fig.9 Stratigraphic correlation and lithofacies distribution characteristics of layers 1-5 of the Dongyuemiao member constrained by short eccentricity in the Fuxing area

但对异地型含介(介壳质)泥岩的岩相组合类型控制明显。

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Lithofacies Distribution and Organic Matter Enrichment of Shale Under the Constraint of Astronomical Cycles: A case study of the Dongyuemiao member in the Fuxing area, Sichuan Basin

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Abstract: [Objective] Significant breakthroughs have been made in the exploration of Jurassic continental shale oil

in the Sichuan Basin. Owing to the small grain size, rapid facies changes, and strong heterogeneity of continental mud shale facies, it is necessary to systematically study the development rules of lithofacies and organic matter enrichment characteristics in the Dongyuemiao member. **[Methods]** Taking the Dongyuemiao member of the Ziliujing Formation in the eastern part of the Sichuan Basin as an example, based on core data, non-core lithofacies interpretation from well logging, and utilizing theory of cyclic stratigraphy, various methods, such as data preprocessing, power spectral analysis, evolutionary spectral analysis, filtering of data, correlation coefficient analysis, and astronomical tuning, were employed to establish a 4-5 level high frequency stratigraphic framework for the study area. Through the sedimentary response of astronomical cycles, the development rules of mud shale lithofacies and organic matter enrichment characteristics under high-frequency chronostratigraphic framework were discussed. **[Results]** (1) The Dongyuemiao member shows a good astronomical cycle signal, and the long eccentricity (405 ka), short eccentricity (128 ka), obliquity (43 ka), and precession (21 ka) astronomical cycles were extracted. There are four long and fifteen short eccentricity cycles. The long and short eccentricity cycles were used as the subdivision units of the fourth-order and fifth-order sequences, establishing a 4-5 level chronostratigraphic framework for the Dongyuemiao member in the study area. (2) Considering rock thins, X-ray diffraction whole-rock analysis, and well logging data, the principle of structure first is employed, with components and sedimentary structures as the primary basis. The Dongyuemiao member is divided into two lithofacies types (indigenous organic-rich laminated mudstone facies and exogenous shell-bearing mudstone and shell limestone facies) and seven lithofacies types (laminated silty-rich argillaceous mudstone facies, massive silt-rich argillaceous mudstone facies, laminated mixed mudstone facies, laminated shell-bearing silty-rich argillaceous mudstone facies, laminated silt-bearing shell-rich argillaceous mudstone facies, laminated silt-bearing clay-rich shelly mudstone facies, massive argillaceous shell limestone facies, and massive argillaceous shell limestone facies). (3) The coupling relationship between paleoclimate, mineral composition, and eccentricity revealed that during periods of high eccentricity and its maximum amplitude, the climate was humid and hot with significant seasonal variations. This led to the input of a large amount of fine-grained material and organic matter from the land and the development of indigenous organic-rich laminated mud shale facies. During periods of low eccentricity and its minimum amplitude, the climate was dry and cold with less input of terrigenous materials. The clay mineral and detrital mineral content were lower, and the lithofacies were dominated by exogenous shell-bearing mudstone facies, which affected the development of organic matter. The Dongyuemiao member is primarily controlled by eccentricity cycles, and the deposition of lacustrine shelly mud shale is controlled by the climate changes driven by the 405 ka and 128 ka eccentricity cycles. **[Conclusions]** Long eccentricity is a key factor controlling the distribution of lithofacies and organic matter enrichment. It controls the ordered development of the indigenous organic-rich laminated mud shale facies and exogenous shell-bearing mudstone and shell-bearing limestone facies. Short eccentricity has a limited impact on organic matter enrichment but significantly controls the lithofacies composition of the exogenous shell-bearing mudstone.

Key words: lacustrine shale oil; lithofacies; lithofacies association; astronomical cycles; organic-matter enrichment; Sichuan Basin