

文章编号:1000-0550(2016)01-0033-16

doi: 10.14027/j.cnki.cjxb.2016.01.003

四川盆地早寒武世龙王庙期沉积特征与古地理

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摘要 基于盆内及其周缘钻孔取芯和露头剖面资料,结合区域背景和龙王庙组地层充填厚度分析认为,四川盆地周缘龙王庙期北邻摩天岭微古陆和汉南古陆、西接康滇古陆,东南和东北分别为威信—黔江、石阡—秀山—龙山和镇巴—巫溪—巴东水下古隆起环抱;盆内呈近东北向展布的川中古隆起和阆中—通江拗陷、江津—奉节拗陷构成“一隆两拗”,南部泸州—筠连—昭通继承性拗陷依然存在;盆地及周缘向东北隅和东南隅分别接入秦岭洋和江南盆地。受此影响,龙王庙期发育连陆碳酸盐岩台地—斜坡—盆地沉积体系,连陆碳酸盐岩台地构成沉积主体,可识别出混积潮坪、局限—蒸发台地、半局限—局限台地和台地边缘沉积相类型;进一步通过碎屑岩含量与陆源方向、石灰岩、白云岩和膏盐岩含量与海域局限性和台地边缘相对位置、颗粒岩含量等分析,阐明了龙王庙期岩相古地理展布样式并建立了沉积模式。结果表明潮坪向海侧、水下古隆起、泸州—筠连—昭通拗陷周缘是颗粒滩有利发育区,乐山和资阳地区经由早寒武世早期的充填拼合作用已形成统一古隆起,震旦纪末期—早寒武世的南北向桐梓—筠连裂陷槽演变为泸州—筠连—昭通台内拗陷,研究成果对拓宽龙王庙组滩控岩溶型储层的勘探领域具有重要的指导作用。

关键词 颗粒滩 古隆起 连陆碳酸盐岩台地 台地边缘 相模式 古地理格局

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中图分类号 P512.2 P534.41 **文献标识码** A

0 引言

近期,四川盆地下寒武统龙王庙组取得重大突破,提交天然气探明储量逾 $4\ 400 \times 10^8\ \text{m}^3$ ^[1-2],证实了四川盆地寒武系优质的油气地质基础条件^[3-4]。早年相关研究因地腹资料限制,只能从小尺度工作中获得启示^[5-8],普遍认为中上扬子地区为典型台地相沉积,上扬子北缘接秦岭洋,东南缘与江南海盆相通。显然这些认识对于盆内精细勘探还远远不够。近年尽管龙王庙组储层研究手段不断更替^[9-10],然而限于现有地球物理勘探技术,针对薄滩型储层预测尚不成熟。还原到储层地质学研究,大型古隆起背景、规模化颗粒滩发育仍是川中龙王庙大型气田形成的关键^[11],颗粒滩叠加加里东期岩溶改造是龙王庙组储层形成的重要地质因素^[12]。可见,弄清大中尺度古地理格局和颗粒滩平面预测依然是寻找盆内有利储层区带的首要工作。

众所周知,颗粒滩平面展布受同期海平面升降和古地理格局控制^[13-14]。对龙王庙期古地理格局的恢复已然成为预测颗粒滩发育的关键。在此基础上,结合沉积相的展布特征,可进一步寻找潜在的颗粒滩发育区和预测规模储集体。随着龙王庙组勘探局面的打开,相关探索性研究已逐步深入^[15-18],但对于龙王庙期古地理格局、沉积相及颗粒滩平面展布的基本认识分歧仍然较大,如蜀南一带元古代—早寒武世存在的裂陷槽是否关闭,大型古隆起对优质储层的形成控制效应^[19-23]具体表现在哪些方面,上扬子台地与秦岭洋、江南海盆连接部为缓坡还是陡坡,颗粒滩相平面展布规律受何种因素控制。对此类问题认识不清,无疑限制了龙王庙期下一步精细勘探。本次在前人基础上,综合曾伟等^[24-25]国家重大专项子课题研究成果,依靠盆地及周缘 171 份钻孔、区域野外露头剖面调查资料(图 1a)进行综合调研、实测踏勘、取样与论证分析,利用地层充填厚度恢复龙王庙期沉积背

收稿日期: 2015-01-19; **收修改稿日期:** 2015-03-25

基金项目: 国家自然科学基金项目(41402126); 国家科技重大专项(2011ZX05004-005-03); 四川省省属高校“天然气地质”科研创新团队建设计划[**Foundation:** National Natural Science Foundation of China, No. 41402126; National Science and Technology Major Project, No. 2011ZX05004-005-03; The Construction Plan For Scientific Research Innovation Team in Provincial Universities for "Natural Gas Geology" in Sichuan Province]

景,综合运用单井和野外露头岩性、岩相资料识别和划分沉积相,并通过单因素定量分析恢复龙王庙期岩相古地理,建立相应沉积相模式,希望研究成果有助于拓宽龙王庙组滩控岩溶型储层勘探领域。

1 区域背景

四川盆地位于扬子地区西北缘(图1b),为两组深大断裂控制形成的菱形构造区。中上扬子地区新元古代主要为西部川滇藏陆块和东部扬子陆块构成(图1c),北部为秦岭—大别裂陷,南部为哀牢山裂陷包围截断,中部成都—昆明一线被龙门康滇裂陷槽南北向贯穿^[26]。受强烈拉张作用^[27]影响,两大陆块逐步下沉,陆壳移离。早寒武世,南北部裂陷带逐步演化为被动大陆边缘坳陷盆地^[28],中部裂陷槽逐渐闭

合,扬子陆块北缘襄樊—广济断裂以北地区^[29]和扬子陆块东南缘江南断裂—武陵、雪峰西侧断裂以东地区^[30]均呈被动大陆边缘盆地性质^[26]。中上扬子地区属康滇古地理体系,发育浅水碳酸盐岩台地相^[8],四川盆地位于康滇古陆东南侧,自西向东发育斜坡和盆地相,构成一套完整的古地理单元序列^[6,31]。龙王庙期川中隆起和黔中隆起已初具雏形(图1c),均属水下古隆起^[32]。根据最新国际上关于寒武系地层4统10阶的划分方案^[33-34],龙王庙期对应于第四阶中上部,与Late Toyonian-Lower Amgan^[31]大致相当。区域资料显示,研究区不同地层小区表现为若干同期异相沉积单元,包括龙王庙组、孔明洞组、清虚洞组和石龙洞组^[35],各组垂向岩相组合和演化差异明显。总体而言,龙王庙期以发育清水碳酸盐岩沉积为主,局

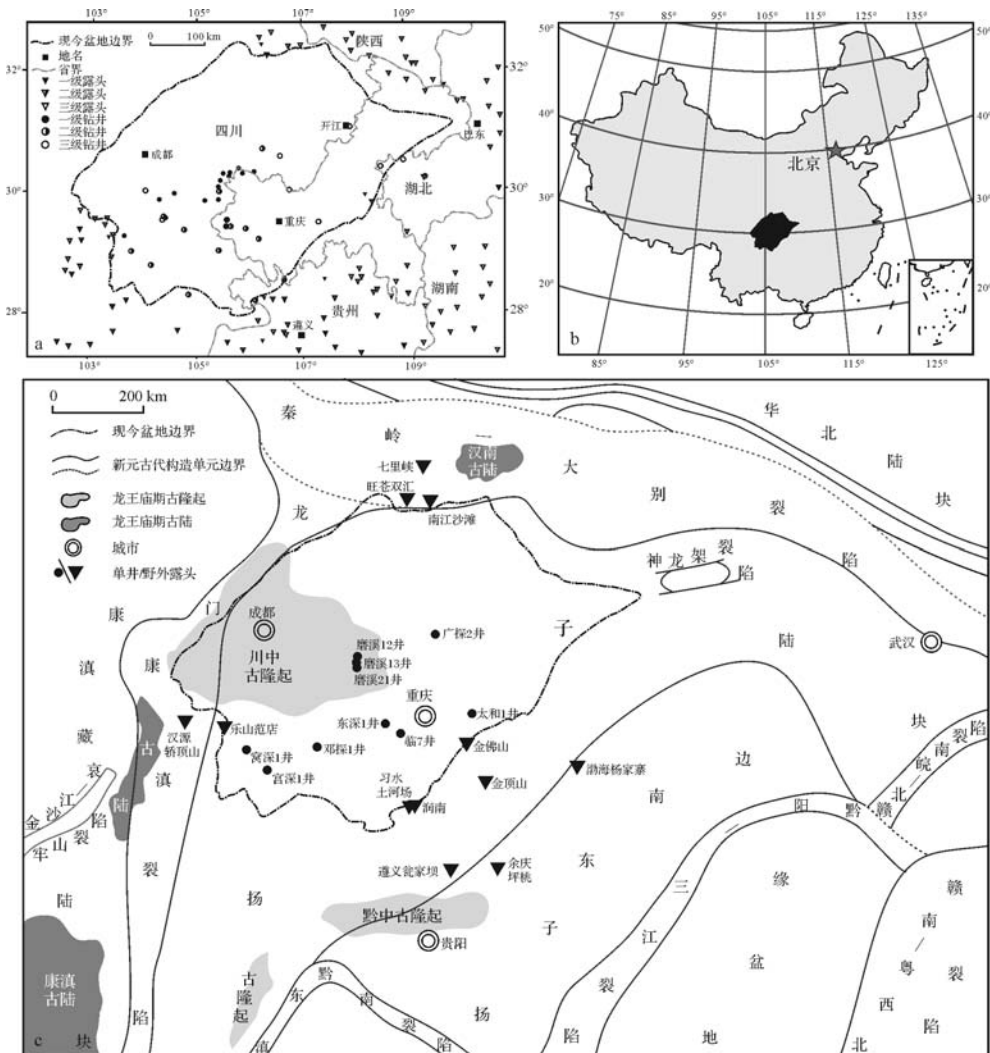


图1 研究区及邻区地质背景

a. 选用单井和野外露头区域位置; b. 研究区所在区域位置; c. 中上扬子地区新元古代早期古构造格局(据周小进等^[26],略有改动)和龙王庙期主要古陆、古隆起位置(据门玉澎等^[38],略有改动)及现今盆地边界和重要单井、野外露头位置

Fig.1 Geological settings of the study area and its adjacent regions

部存在典型蒸发岩相,陆源碎屑供应相对不足。

2 龙王庙期沉积背景

综合龙王庙期地层古生物、岩性及电性特征,对各单井和野外剖面地层厚度进行统一梳理统计。如图2,盆地西部北到广元,中抵资阳—绵阳,南达石棉—冕宁—一线以西地区遭受广泛剥蚀,克拉通盆地内部均为连续沉积,呈西北薄东南厚、北东南西向厚薄交互展布,区内地层厚度多介于60~220 m,其中川南泸州—筠连—昭通,川东重庆、忠县—云阳厚度普遍较高,多介于340~380 m,核部大于380 m,盆外东北和东南部各存在一处厚度高值区,东北部城口北—镇坪—巫溪—兴山一带,普遍介于180~220 m,东南部铜仁西—花垣—大庸南一般在300 m以上,核部大于380 m。

前人研究表明,龙王庙期经历了两个完整三级海侵海退旋回的二级海侵^[36],海水沉积能量恢复对填平补齐作用至关重要。高能海水淘洗和细粒沉积物向低地回流作用致使古海底地势较低地区地层厚度大于地势较高地区。据此,可将台内厚度较小地区划为水下隆起区,厚度较大地区拟定为坳陷区。如图2,川中古隆起所在乐山和资阳一带地层厚度通常在60~100 m之间,厚度呈西浅东深且较稳定,明显小于非古隆起区的100~300 m,由此可推断龙王庙期川中地区经由早寒武世早期的充填拼合作用已形成统一的古隆起,上覆于刚性基底之上的古隆起沉积演化

是地层厚度相对较小的直接原因。在一个二级海侵期,古隆起无疑为优质颗粒滩发育提供了较好水下高能沉积场所。

南缘泸州—筠连—昭通一带仍然存在南北走向地层厚度异常增厚,这与寒武纪初期贯穿上扬子地区的张性裂陷槽未完全闭合^[37]是一致的。邓探1井钻遇龙王庙组399 m,普遍高于周缘100~200 m,以石灰岩和石膏互层,夹少量白云岩、云质灰岩和灰质云岩的岩性组合特征也进一步印证了这一认识,此时裂陷槽已由深水海槽沉积转换为浅水台地坳陷区,同理可推测川中古隆起北部钻孔和野外露头稀缺的阆中—通江一带极可能与南部凹陷具有类似的厚度展布特征。然而,受钻孔资料钻遇断层影响,很难就川东重庆一带地层厚度进行统计恢复。结合四川盆地普遍存在的蒸发岩相特征^[38-40],川东一带东深1井和临7井普遍存在的大套膏盐岩证实了川东海域水体总体受限,盐度较高,推测其最大地层厚度大于380 m,太和1井钻遇龙王庙组151 m,泥—粉晶白云岩夹石膏岩的垂向特征也进一步印证了海水受限坳陷带的存在。

结合龙王庙期区域背景和地层厚度平面展布,可认为四川盆地北邻摩天岭微古陆^[41]和汉南古陆^[42]、西接康滇古陆^[43],东南和东北分别为威信—黔江、石阡—秀山—龙山和镇巴—巫溪—巴东水下古隆起环抱;盆内呈近北东向展布的川中古隆起和阆中—通江坳陷、江津—奉节坳陷构成的“一隆两坳”,盆地向东

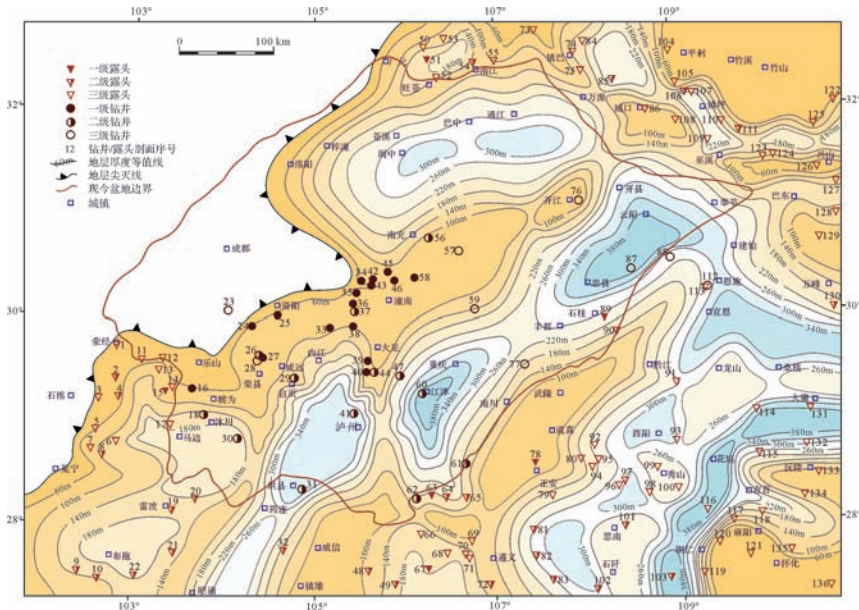


图2 中上扬子地区龙王庙期地层厚度等值线图(图中序号所代表钻井和野外露头名称如表1所示)

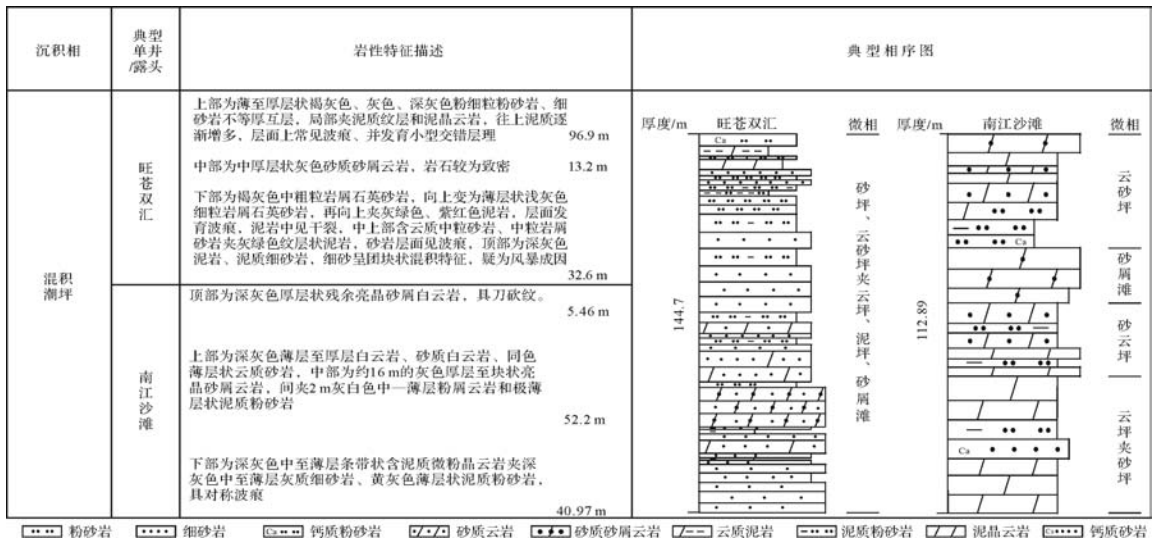
Fig.2 Contour map of thickness during Longwangmiao period in Middle-Upper Yangtze region

(numbers represent wells and outcrops seen in Table 1)

表1 四川盆地及邻区所用钻井和野外露头一览

Table 1 Used wells and outcrops in the Sichuan Basin and its adjacent regions

单井(露头)序号	位置	级别	单井(露头)序号	位置	级别	单井(露头)序号	位置	级别	单井(露头)序号	位置	级别
1	白井沟	3	35	安平1井	1	69	庙子湾	3	103	岑巩羊桥	2
2	汉源轿顶山	2	36	高科1井	1	70	花尖山	3	104	松树埡	3
3	林罗沟	3	37	高石1井	2	71	台古山	3	105	岚皋支河	3
4	田坪	2	38	高石6井	1	72	遵义瓮家坝	2	106	江西街	2
5	白沙沟	3	39	荷深1井	1	73	汉南	3	107	龙门桥	3
6	甘洛波波乡	3	40	螺观1井	1	74	高桥	3	108	东安	3
7	敏子洛木	3	41	阳深2井	2	75	九拱桥	3	109	镇坪大雄溪	3
8	越西碧鸡山	2	42	磨溪204井	1	76	五科1井	3	110	镇坪	3
9	大槽河	2	43	磨溪17井	1	77	太和1井	3	111	竹溪柳林店	2
10	洛乌沟	2	44	盘1井	2	78	金顶山	1	112	鄂参1井	3
11	张村	3	45	磨溪202井	1	79	茶林	3	113	恩施白果坝	3
12	六道河	3	46	宝龙1井	1	80	三坑	3	114	龙山砂坝	3
13	峨眉张山	3	47	东深1井	2	81	坨田坝	2	115	波罗寨	3
14	潮水溪	3	48	毕节播扎	2	82	十里溪	2	116	三宝	3
15	乐山范店	1	49	大方施梨	2	83	余庆坪桃	2	117	凤凰七梁	3
16	老龙1井	1	50	旺苍母家沟	3	84	镇巴兴隆场	3	118	泸溪兴隆场	3
17	雪口山	3	51	双汇正源	1	85	万源蒲家	2	119	贵州万山	3
18	窝深1井	2	52	天台	3	86	城口石溪河	3	120	凤凰十八坪	3
19	雷波抓崖	2	53	青木洞	3	87	建深1井	3	121	麻阳江口	3
20	永善长坪	2	54	南江沙滩	2	88	利1井	3	122	大红坊	3
21	永善金沙	2	55	贵民关	3	89	石柱六塘	1	123	艾五坪	3
22	对坪	2	56	南充1井	2	90	石柱	2	124	小当阳	3
23	油1井	3	57	广探2井	3	91	酉阳小咸	3	125	龙头沟	3
24	资2井	1	58	女基井	1	92	石界水	3	126	南阳河	3
25	资4井	1	59	座3井	3	93	渤海杨家寨	3	127	建阳坪	3
26	威寒26井	2	60	临7井	2	94	泥矿山	3	128	庙河	3
27	威寒1井	1	61	丁山1井	2	95	三角木	3	129	长阳两河口	3
28	威寒101井	3	62	林1井	2	96	毛田	3	130	杨家坪	2
29	自深1井	2	63	习水土河场	1	97	平井	3	131	大庸田坪	3
30	官深1井	2	64	润南	2	98	木盆溪	3	132	沅陵王家坪	3
31	宁2井	2	65	九坝	2	99	秀山溶溪	3	133	沅陵凉水井	3
32	镇雄罗坎	2	66	马跃水	3	100	邓阳坳	3	134	田家坪	3
33	高石17井	1	67	金沙岩孔	1	101	印江后坝	2	135	怀化花桥	3
34	磨溪12井	1	68	石塔	3	102	石阡窑上	2	136	隆回大水田	3



注:南江沙滩野外剖面引自参考文献[44]

图3 中上扬子地区龙王庙期混积潮坪典型垂向岩性组合与相序图(剖面位置见图1)

Fig.3 The representative vertical lithological assemblage and sequence of mixed tidal flat during Longwangmiao Formation, Middle-Upper Yangtze region(well/outcrop locations seen in Fig.1)

北和东南分别接入秦岭洋和江南盆地。

3 龙王庙期沉积相类型及特征

如前文所述及,龙王庙期发育台地—斜坡—盆地沉积体系,台地相构成四川盆地沉积主体。根据野外露头观察、钻孔取芯和测录井资料,共识别划分出混积潮坪、局限—蒸发台地、半局限—局限台地,综合前人研究成果,台地相远端还存在台地边缘相带。

3.1 混积潮坪

常见于盆地西南缘荣经—石棉一线以东、资阳—乐山一线,以及盆地西北缘的广元—旺苍—南江—镇巴—紫阳一线。根据沉积物质和特征的差异,可进一步识别出潮上带、潮间带和潮上带亚相及多种微相类型,包括潮道、潮缘滩、砂坪、云坪、灰坪、泥坪、膏坪及多种混合坪等。按照陆源碎屑与碳酸盐岩含量和叠置方式的不同,共识别出两类混积潮坪类型,一类以旺苍双汇(图3)为特征的陆源碎屑夹碳酸盐岩混积潮坪,主要为灰色石英岩屑砂岩夹薄层弱还原—氧化色的泥岩、灰色砂质砂屑云岩,发育典型的波痕(图4a)、潮道(图4b)、平行层理、交错层理(图4c)等指示复杂水动能环境的沉积构造。另一类以南江沙滩(图3)较为典型,发育碳酸盐岩夹陆源碎屑垂向序列,主要为深灰色云岩、砂质云岩、砂屑云岩夹薄层状泥质粉砂岩、云质砂岩,自下而上砂质成分逐渐增多,发育脉状层理、水平层理,底部粉砂岩中发育的对称波痕指示极浅水动能环境。

3.2 局限—蒸发台地

常见于盆地西南缘布拖—马边、盆地中部江津—南川一带,以发育大量准同生白云岩、蒸发岩类为特征,根据沉积物质和特征的不同共识别出若干类亚相类型,包括局限潟湖、蒸发潟湖、台内滩、滩间海、风暴岩亚相(图5)。常见灰—深灰色泥—粉晶云岩的局限潟湖夹白色石膏、石盐矿物的蒸发潟湖和灰—深灰色颗粒白云岩的颗粒滩,灰—深灰色颗粒白云岩的台内滩亚相夹同色泥晶云岩的滩间海亚相垂向序列。泥—粉晶云岩类常见水平层理,颗粒岩中以砂屑、鲕粒为主(图4d),次以核形石、生屑、藻屑常见,发育交错层理、冲刷面、正粒序(图4d)和少量逆粒序等。蒸发矿物常呈自形或团块状(图4e),常见膏模孔,常夹风暴序列。盆外常见岩溶角砾岩(图4f),盆内钻孔取芯录井(如临7井)中见大套膏盐岩沉积。风暴岩序列通常小于20 cm,表现为中—薄层状颗粒云岩与低能正常潟湖沉积的泥粉晶云岩、泥云岩、含膏泥粉

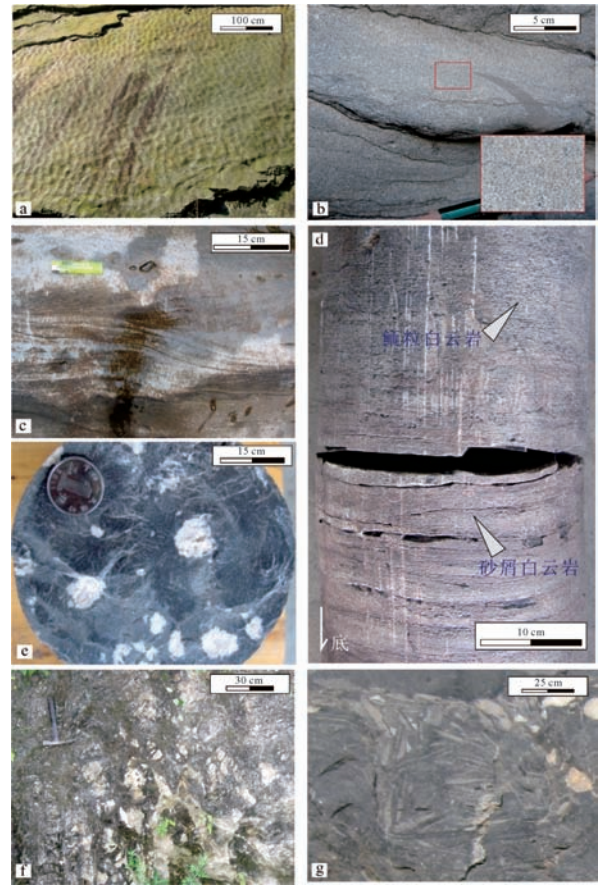


图4 中上扬子地区龙王庙期典型沉积构造与岩性特征 a.波痕(野外剖面,七里峡,龙王庙中上部);b.潮汐水道中的砂质鲕粒云岩(野外剖面,七里峡,龙王庙中上部);c.浪成交错层理(野外剖面,七里峡,龙王庙中上部);d.向上变粗的颗粒滩旋回,逆粒序层理发育(磨溪21井,4 661.78~4 662.11 m);e.泥粉晶云岩中的石膏团块,现已为白云石充填(单偏光薄片,磨溪13井,4 488.73 m);f.岩溶角砾岩(野外剖面,金佛山,龙王庙组中上部);g.风暴潮破碎的竹叶状砾屑灰岩,下部为砂屑灰岩(野外剖面,金顶山,龙王庙组中上部)。

Fig.4 Sedimentary structures and lithological features during Longwangmiao period in Middle-Upper Yangtze region

晶云岩、膏质泥粉晶云岩,底部具有明显底冲刷面,向上为正粒粒序递变,砾石无定向杂乱排列。

3.3 半局限—局限台地

主要分布于昭通—筠连—遵义—正安—道真—黔江—建始,受古隆高地水体包绕,海水循环局部受限(图6)。整体发育灰色泥晶(灰)云岩和泥质(灰)云岩、泥晶云岩夹颗粒灰(云)岩和风暴成因薄层状的砂(砾)屑(灰)云岩的风暴岩(图4g),底部常见灰—青灰色钙质砂岩、钙质页岩、灰—深灰色泥灰岩、鲕粒灰岩夹层,下部泥质含量较高,常见灰色厚层豹皮状、条带状灰岩夹层,黔北地区石阡一带龙王庙组

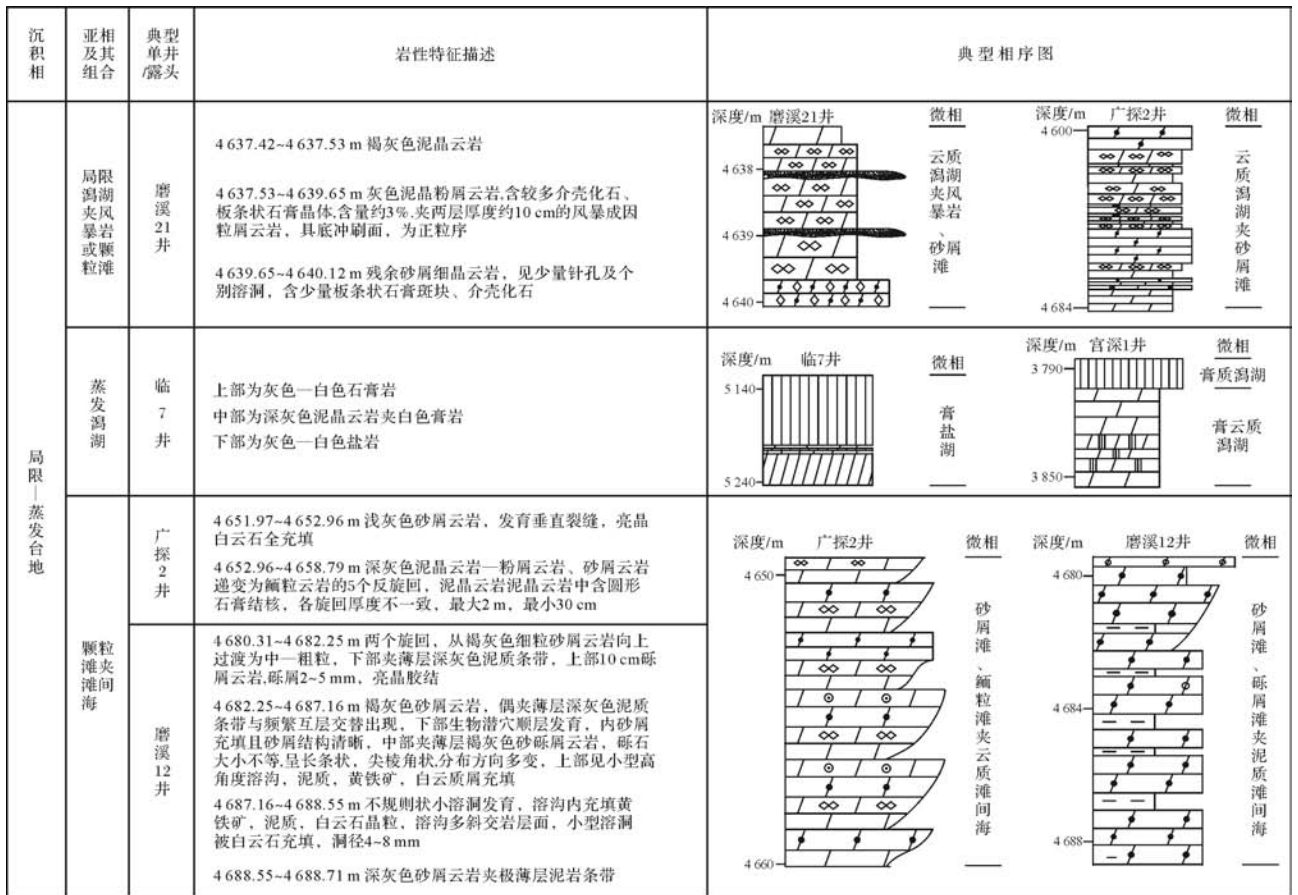


图5 中上扬子地区龙王庙期局限—蒸发台地典型垂向岩性组合与相序图(剖面位置见图1)

Fig.5 The representative vertical lithological assemblage and sequence of restricted to evaporate platform during Longwangmiao Formation, Middle-Upper Yangtze region(well/outcrop locations seen in Fig.1)

中上部见约5 m薄层状灰绿色粉砂—细砂岩,上部厚层状细晶—中晶白云岩夹层较多,局部地区顶部逐渐变为灰黄色、灰—深灰色砂质、泥质白云岩。半局限—局限台地具有与局限—蒸发台地具有较为类似的沉积构造,不同之处在于随蒸发岩相的缺失,水体逐渐开阔,向东和东南方逐渐由白云岩向石灰岩类过度。

3.4 关于台地边缘的存在

大量勘探实践表明,台缘滩可形成颇具规模的优质储集体^[45-49]。龙王庙期扬子地区典型台地边缘相带的存在问题一直饱受争议^[15,50-51]。镶边台地模式和缓坡台地模式的不同常表现为沿斜坡带是否出现规模连片的礁滩相。早年郑荣才和曾允孚^[52]于湘西渔塘地区发现了厚约120~160 m具有相当规模的堤状藻礁,生物礁明显受铜仁—大庸深大断裂带的控制。近年认为,深大断裂活动是黔东—湘西地区大量高能颗粒滩、生物礁丘相形成的重要地质因素,铜庸

深大断裂西侧上升盘的存在是镶边台地边缘存在的直接证据^[32];另有指出,湘西—黔东地区早期为浅海陆棚—潮下高能带—潮间带旋回的缓坡特征,晚期经历缓坡向浅滩—潮间带旋回的镶边台地边缘过度^[53]。相比而言,针对上扬子北缘与南秦岭洋构造带之间龙王庙组研究相对较少;相关报道认为,该区台缘滩陡坡沉积特征初现端倪^[54-56],晚寒武世陡坡特征更为成熟^[56]。就本次地层厚度统计结果(图2)来看,东北部城口—兴山和东南部铜仁西—花垣—大庸南都存在地层厚度异常,分别指示扬子北缘地区南秦岭大陆边缘裂谷带同沉积断裂^[55]和新元古代雪峰—四堡岛弧造山带活动^[57],二者厚度异常增厚可能与斜坡进积体与滑塌作用相关,这也可从广泛发育的暗色石灰岩矿物相、滑揉构造和垮塌等重力流沉积构造得到佐证^[6,55,58-61]。如前文所述,东北部和东南部较为相似的古构造背景,暗示二者极可能具有类似沉积组合,均发育潜在规模的台地边缘。

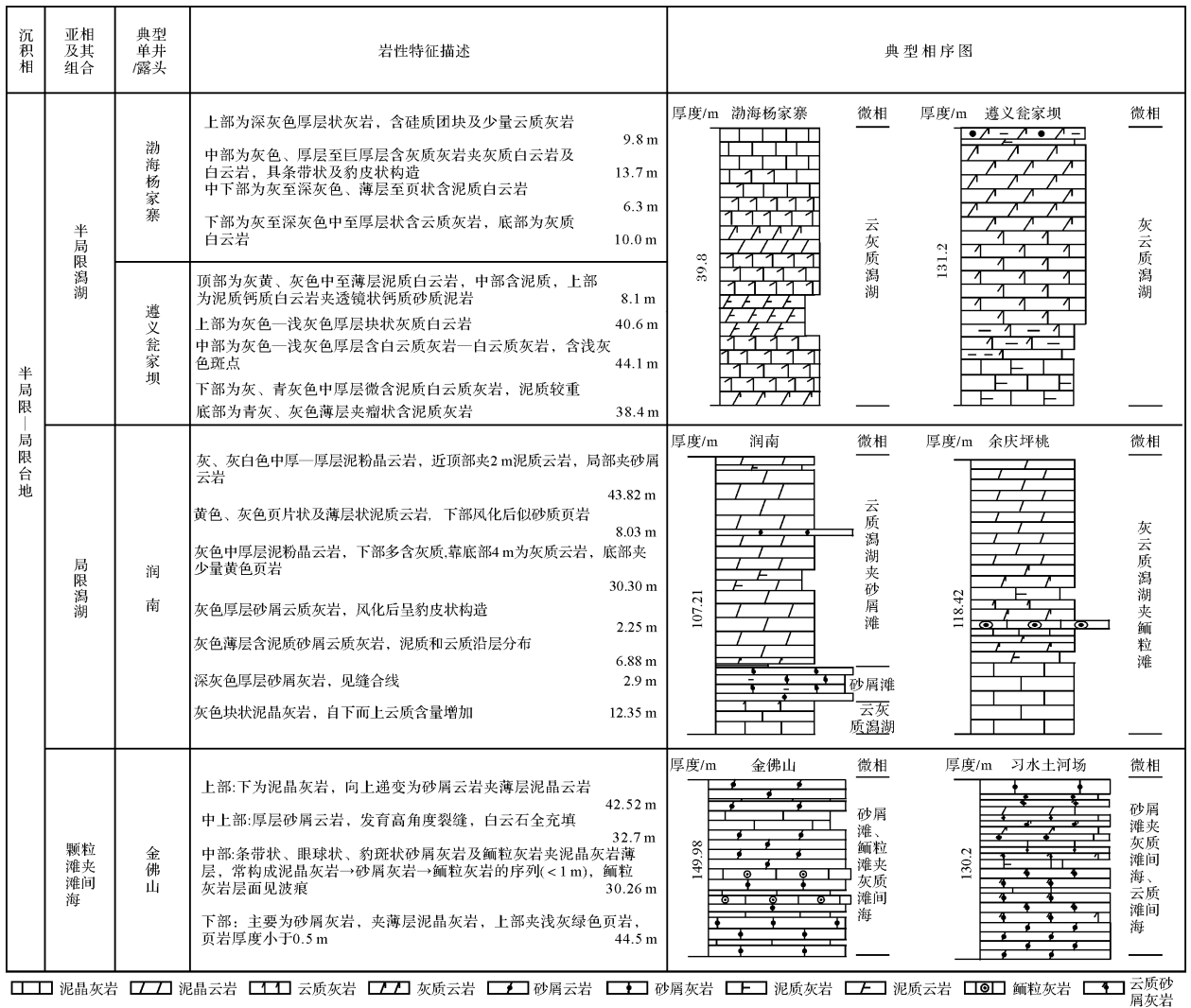


图6 中上扬子地区龙王庙期半局限—局限台地典型垂向岩性组合与相序图(剖面位置见图1)

Fig.6 The representative vertical lithological assemblage and sequence of semi-restricted to restricted platform during Longwangmiao Formation, Middle-Upper Yangtze region (well/outcrop locations seen in Fig.1)

4 岩相古地理

基于野外剖面实地测量、岩芯描述、岩屑录井和测井解释成果,对85口单井和野外剖面点进行岩性分类统计,编绘单因素分析图件,包括浅水陆源碎屑岩厚度/地层厚度等值线图(图7)、石灰岩厚度/地层厚度等值线图(图8)、白云岩厚度/地层厚度等值线图(图9)、膏盐岩厚度/地层厚度等值线图(图10)、颗粒岩厚度/地层厚度等值线图(图11)。

陆源碎屑含量(图7)包括从陆源区搬运至台地内沉积的浅水砾岩、砂岩、粉砂岩和黏土岩厚度之和与地层总厚的厚度百分比。如图7,碎屑岩集中分布于剥蚀线附近西北缘的广元—旺苍—南江和西南缘

的石棉东—荣经两处海域。西北缘陆源物质影响范围相对较高,粒度普遍较大,多为中砂—细砂级,自西向东砂质含量逐渐降低,粉砂质、泥质逐渐增高。西南缘陆源物质影响范围明显偏小,以粉砂和泥级为主,至汉源轿顶山一带陆源碎屑与碳酸盐岩混积特征逐渐明显,至窝深1井碎屑岩含量仅12%。海陆过度区的陆源物质展布形态与北东南西向的剥蚀线基本一致,并与地层厚度等值线展布近似相等,表明陆源物质主要来自于西北部摩天岭微古陆,西南部康滇古陆和北部汉南古陆影响相对较弱。

图8表明,石灰岩主要集中于盆地外缘,其中西南部昭通—筠连—泸州、石棉南,东北部镇巴东—镇坪—兴山北、巴东西和东南部道真—黔江—宣恩—五

峰南,铜仁—吉首—沅陵较为富集,盆内除绵阳—南江一带有少量发育外,石灰岩含量均低于10%。相较而言,白云岩则集中发育于盆地内部(图9),除马边—泸州—内江、重庆、忠县—云阳和阆中—通江一带外,含量普遍大于80%,盆地外缘西南部布拖—雷波—冕宁南、南部遵义南—正安南,石阡东—秀山—

大庸、麻阳和沅陵呈片状展布,北部南江—镇巴、五峰—建始—兴山分别与盆地北部、东部发育区相接。膏盐岩含量百分比(图10)显示,川东重庆—南川一带最为发育,川北阆中—通江、川东开县—忠县和川南马边、布拖较为发育,东缘巴东东部发育规模相对较小。石灰岩、白云岩和膏盐岩平面展布及含量叠加

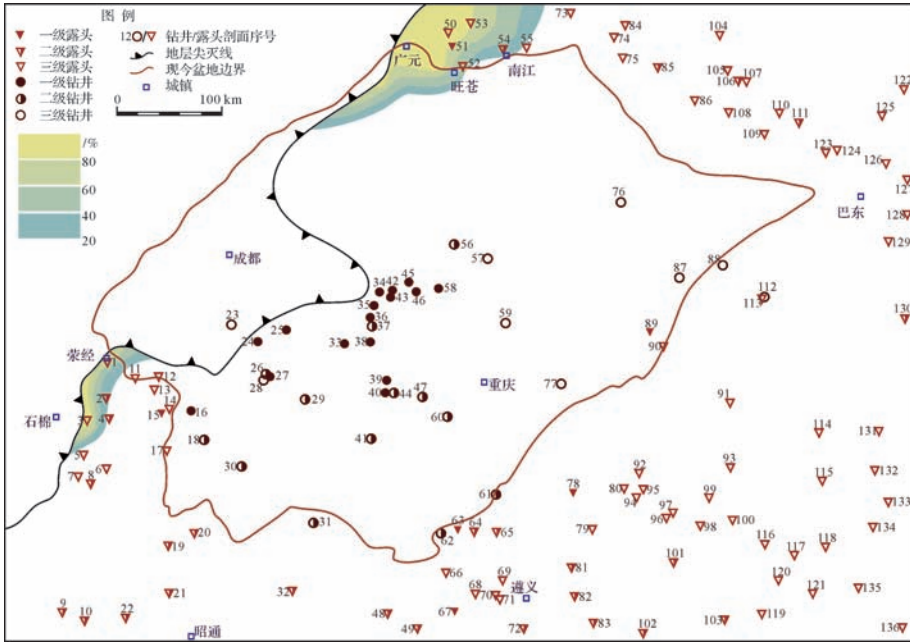


图7 中上扬子地区龙王庙期陆源碎屑岩厚度/地层厚度等值线图

Fig.7 The thickness content contour map of terrigenous classic rocks to the gross thickness during Longwangmiao Formation, Middle-Upper Yangtze region

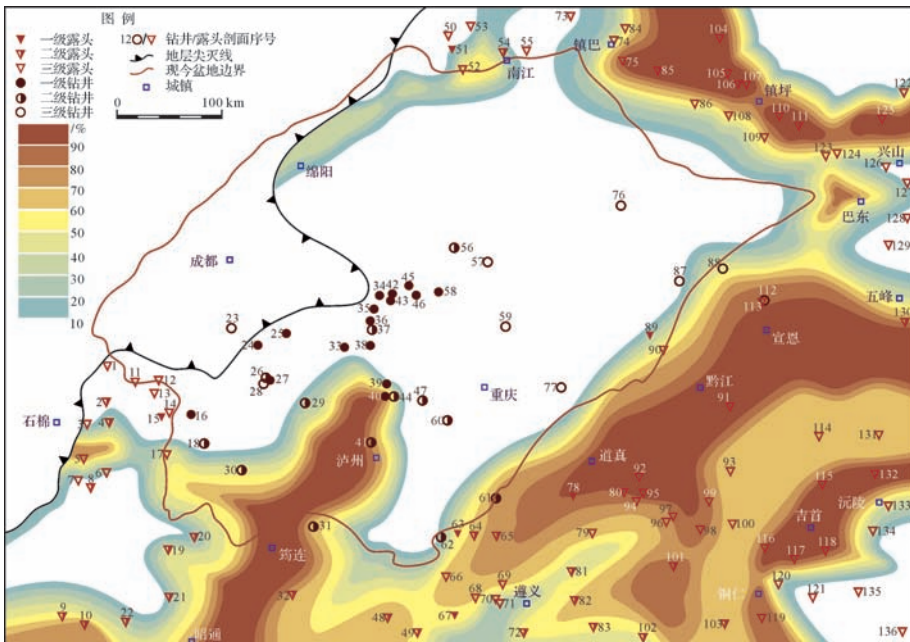


图8 中上扬子地区龙王庙期石灰岩厚度/地层厚度等值线图

Fig.8 The thickness content contour map of limestones to the gross thickness during Longwangmiao Formation, Middle-Upper Yangtze region

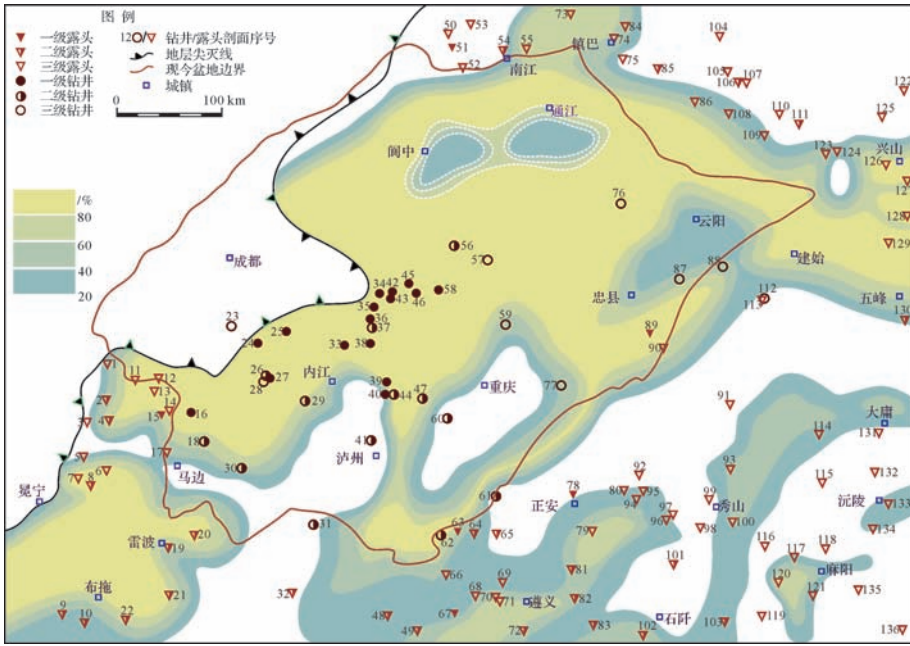


图9 中上扬子地区龙王庙期白云岩厚度/地层厚度等值线图

Fig.9 The thickness content contour map of dolostones to the gross thickness during Longwangmiao Formation, Middle-Upper Yangtze region

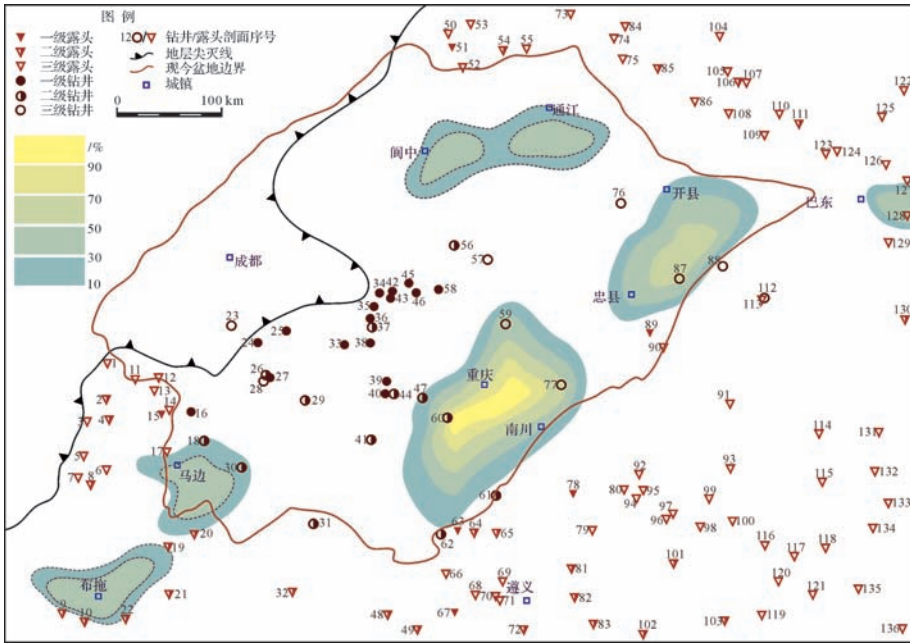


图10 中上扬子地区龙王庙期膏盐岩厚度/地层厚度等值线图

Fig.10 The thickness content contour map of gypsum and halite rocks to the gross thickness during Longwangmiao Formation, Middle-Upper Yangtze region

关系表明,盆内以发育局限—蒸发环境的云岩类和膏盐岩类沉积^[38,40],盆地外围逐步向石灰岩沉积过度,水体逐渐开阔、盐度趋于正常,指示半局限—局限相。

盆外受诸多因素影响,三大岩类百分比关系存在差异,西北缘和西南缘主要受陆源注入供给影响以陆源碎屑砂和石灰岩类沉积为主;南缘泸州—筠连—昭通一带则受裂陷槽基底影响,海水持续循环冲刷,以

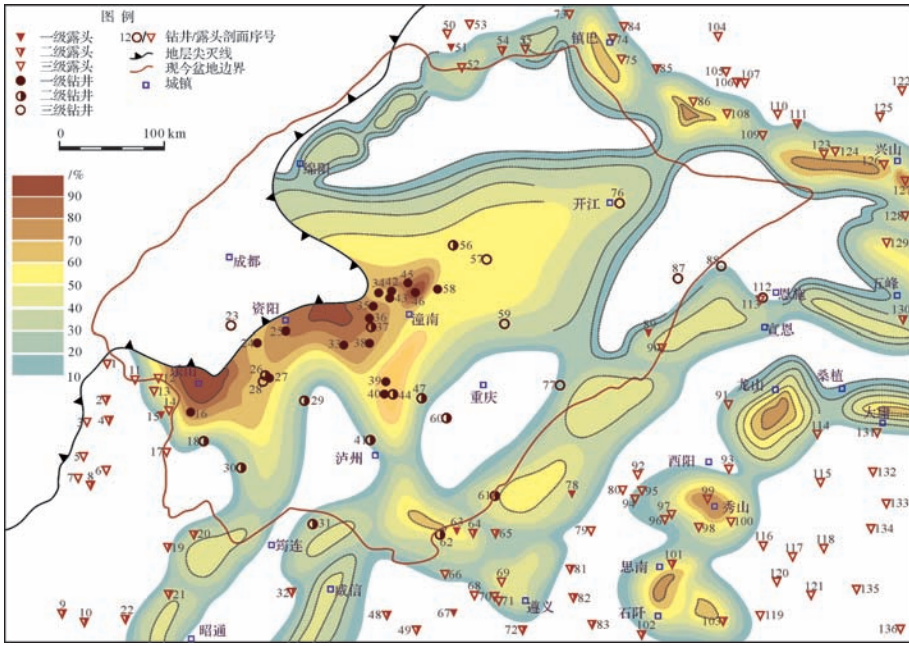


图 11 中上扬子地区龙王庙期颗粒岩厚度/地层厚度等值线图

Fig.11 The thickness content contour map of shallow grain rocks to the gross thickness during Longwangmiao Formation, Middle-Upper Yangtze region

石灰岩沉积为主;东南缘和东北缘则因继承性基底断裂带的存在,水体循环局部受限,白云岩和石灰岩交互占据。颗粒结构碳酸盐岩常指示高能界面的存在^[13],高能界面受基底底形引起的海水水动能控制^[61]。前文述及,上扬子地区龙王庙期极可能存在台地边缘高能相带。研究区东南隅和东北隅浅水岩类欠发育,暗示其向陆侧东南缘石阡—思南—龙山—桑植和东北缘镇巴—兴山为台地边缘高能相带。以此为指导,结合古地理背景中水下古隆起相对位置,最终完成颗粒岩含量等值线图。如图 11,北部绵阳—镇巴北所处潮坪向海侧、乐山—开江所在川中古隆起,川中古隆起东南侧北东向的遵义—恩施—五峰南水下古隆起、泸州—筠连—昭通拗陷周缘一带为台内颗粒滩有利发育区,东南缘石阡—思南—龙山—桑植和东北缘镇巴—兴山一带为颗粒滩发育潜在区域。南缘泸州—筠连—昭通一带由于台内拗陷的存在,致使拗陷区内部为低能沉积,拗陷区东西两侧颗粒滩较发育。

基于上述五幅单因素图件分析,参照前人定量划分标准^[6-7,15,62],针对四川盆地龙王庙期各单因素进行了定量划分(表 2),结合沉积相展布规律定性的研究认识,以及区域其他环境判别标志,编绘中上扬子地区龙王庙期的岩相古地理平面展布图(图 12)。如

图 12,四川盆地及其周缘所在的中上扬子地区,西部冕宁—石棉东—荣经—乐山—资阳—绵阳—广元以西为古陆发育区,冕宁—乐山西和绵阳—南江地区分别过渡为混积潮坪海陆过度相。向东进入台地沉积,乐山—开江所在川中古隆起带发育相当规模的巨型浅滩化台内滩,川中古隆起西北侧的苍溪—通江—开县和东南侧的泸州—丰都地区,水体闭塞,发育大套膏盐岩为特征的局限—蒸发潟湖亚相,二者构成蒸发—局限台地主体。川中古隆起东南侧的威信—珙县—正安—武隆—石柱—恩施东一线,为古隆起东南侧近平行排列的水下高地,发育一系列北东南西向的台内滩带,其西南部自贡—镇雄—遵义—正安—道真—黔江—宣恩—云阳—巴东“V”型海域,为水体局部受限的半局限潟湖亚相,二者构成局限—半局限台地主体。直至东北部的镇巴—万源—城口—巫溪—兴山和东南部的石阡—思南—龙山—桑植,相变为台缘滩和台缘滩间海的环带状台缘带,分别向东北部镇巴东—镇坪—兴山北和东南部的铜仁—大庸南相变为灰质和云灰质半深海斜坡。东北隅的平利—竹溪—竹山和东南隅怀化分别过度为秦岭洋和江南盆地,从炭质板岩、千枚岩等轻微变质岩性可以推测在沉积期,其岩性多为炭质泥和泥质沉积。

国内外学者曾先后提出了多种碳酸盐岩沉积模

表 2 单因素分析定量叠加标准一览

Table 2 The quantitative superposition standard for single factor analysis

沉积相划分			岩性分类统计 (含量/%)					
相	亚相	微相	陆源碎屑岩	石灰岩	白云岩	颗粒岩	膏盐岩	
连陆碳酸盐台地相	混积潮坪		>20	>10		<30		
		潮缘滩		>20	>10		>30	
	局限—蒸发台地	蒸发潟湖	含膏潟湖				<30	10~30
			膏质潟湖				<30	30~50
			膏盐湖				<30	>50
	半局限—局限台地	局限潟湖			<10	>20	<30	<10
		台内滩					>30	
	台缘带 (斜坡进积体向陆侧)	半局限潟湖			>10		<30	
		台缘滩间海					10~30	
		台缘滩					>30	
斜坡相 (斜坡进积体所在区)			台缘带外侧至石灰岩含量>50%的地区					
盆地相 (斜坡进积体广海侧)			斜坡带外侧石灰岩含量<50%的地区					

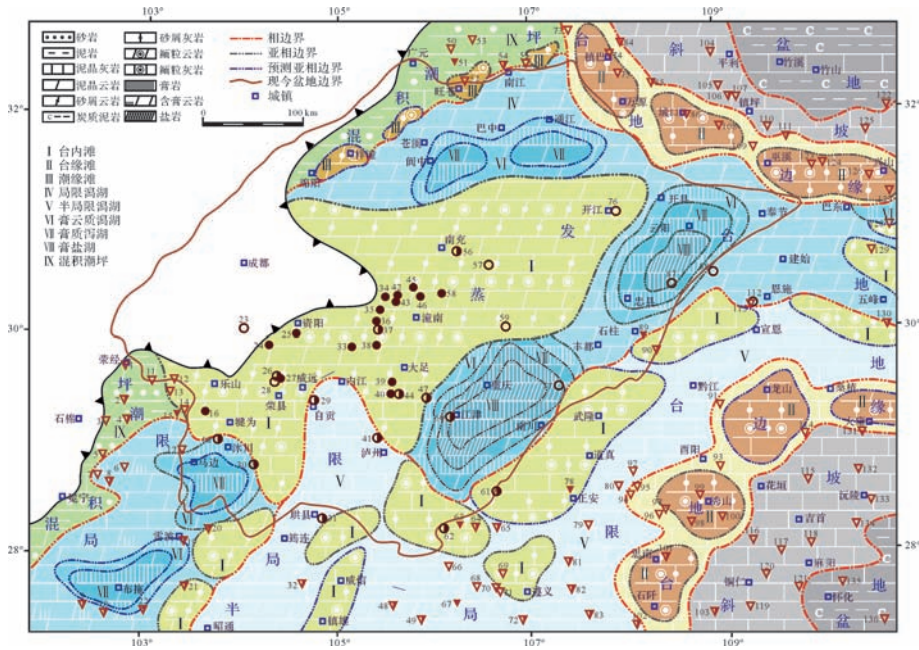


图 12 中上扬子地区龙王庙期岩相古地理平面展布图

Fig.12 The distribution pattern of lithofacies paleogeography during Longwangmiao Formation, Middle-Upper Yangtze region

式^[63-65],目前得到广泛采用的是斜坡—台地沉积模式。综合龙王庙期沉积背景和岩沉积相带时空配置关系,总的说来,西部康滇古陆并不活跃,北部受低缓汉南古陆^[56]的影响亦十分微弱,仅西北部受摩天岭古陆局部影响,由陆向海发育碳酸盐岩与细粒陆源碎屑岩的混积潮坪沉积,盆内仍以清水碳酸盐岩台地建

造为主,具有典型台地—陆表海沉积相模式(图 13)。川中古隆起一带以沉积颗粒云岩的颗粒滩夹泥粉晶云岩的局限潟湖为特色,向东和东南方向逐渐过度为泥晶云岩、膏质云岩(云质膏岩)、膏岩和盐岩的局限—蒸发潟湖夹颗粒白云岩的颗粒滩,构成一类盐度较高,水体循环不畅的局限—蒸发台地相,其中川北

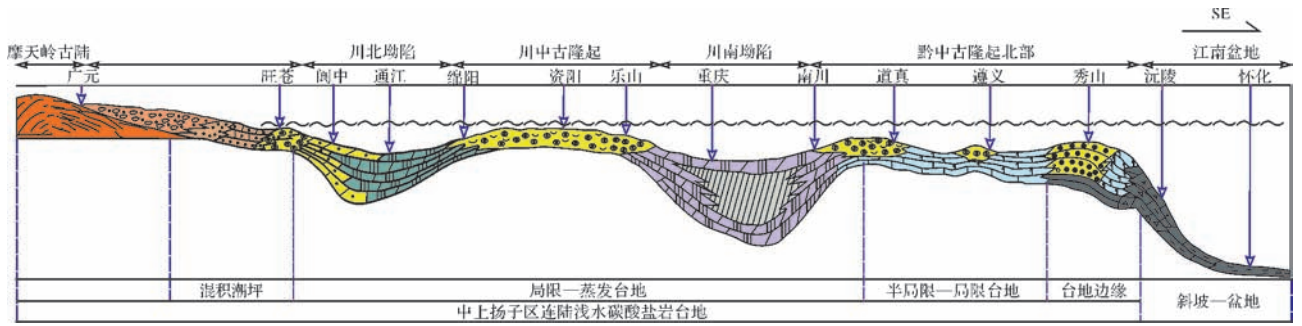


图 13 中上扬子地区龙王庙期沉积相发育模式

Fig.13 The sedimentary model during Longwangmiao Formation, Middle-Upper Yangtze region

和川东坳陷区构成典型的蒸发岩相区,后者蒸发海域特征更强;再向东和南部灰岩逐渐增多增厚,蒸发岩类欠发育,水体趋于通畅,盐度逐渐正常,连同黔中古隆起北部地区在内的广大地区以泥—粉晶石灰岩、泥—粉晶白云岩和泥粉晶灰质云岩(云质灰岩)的半局限—局限潟湖夹颗粒白云岩(石灰岩)、颗粒云质灰岩(灰质云岩)的颗粒滩为特色的半局限—局限台地;东南部和东北部则进入高能台缘带,并向两个方向进一步发育半深海斜坡和深海盆地相。绵阳—镇巴北、乐山—开江、遵义—恩施—五峰南、泸州—筠连—昭通,石阡—思南—龙山—桑植和镇巴—兴山可作为颗粒滩发育的有利区带和有利储层发育区。

5 结论

(1) 四川盆地所在的中上扬子地区其周缘三面为古陆包绕,北邻摩天岭微古陆和汉南古陆、西接康滇古陆,东南和东北分别为威信—黔江、石阡—秀山—龙山和镇巴—巫溪—巴东水下古隆起环抱。盆内呈近北东向展布的川中古隆起和阆中—通江坳陷、江津—奉节坳陷构成的“一隆两坳”,乐山和磨溪地区经由早寒武世早期的充填拼合作用已形成统一的古隆起,震旦纪末期—早寒武世的南北向桐梓—筠连裂陷槽演变为泸州—筠连—昭通天内坳陷。

(2) 陆源碎屑含量表明,西北部摩天岭微古陆为主要陆源区,西南部康滇古陆和汉南古陆影响相对较小;石灰岩、白云岩和膏盐岩含量表明,盆内以发育局限—蒸发环境的云岩类和膏盐岩类沉积,盆地外围逐步向石灰岩沉积过度为主,间夹白云岩沉积,水体具有逐渐开阔、盐度逐渐正常的趋势。颗粒岩含量表明,绵阳—镇巴北所处的湖坪向海侧、乐山—开江所在的川中古隆起,威信—黔江、石阡—秀山—龙山和

镇巴—巫溪—巴东水下古隆起、泸州—筠连—昭通坳陷周缘为台内颗粒滩有利发育区,东南缘石阡—思南—龙山—桑植和东北缘镇巴—兴山一带为颗粒滩潜在区域。

(3) 中上扬子地区具有典型台地—陆表海沉积相模式,发育连陆碳酸盐岩台地—斜坡—盆地沉积体系,川中地区为局限—蒸发潟湖夹颗粒滩的局限—蒸发台地相,向东和南部过渡为半局限—局限潟湖夹颗粒滩相为主的半局限—局限台地相,东南部和东北部进入高能台缘带,向两个方向进一步发育半深海斜坡和深海盆地相。

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Sedimentary Characteristics and Lithofacies Palaeogeography during Longwangmiao Period of Early Cambrian, Sichuan Basin

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Abstract: In order to improve the understanding on the sedimentary characteristics and lithofacies palaeogeography in the Lower Cambrian Longwangmiao Formation of the Sichuan Basin and its adjacent regions, a relative comprehensive analysis on the boreholes and outcrops data were performed. Particularly, by examining the regional geological settings and the thickness distribution pattern within Longwangmiao Formation, this paper found that the Sichuan Basin and its adjacent areas are bordered by Motianling Micro-land and Hannan Oldland in the north, Kangdian Oldland in the west. In the southeast and northeast regions within the study area, were embraced by underwater palaeouplift named Weixin-Qianjiang, Shiqian-Xiushan-Longshan and Zhenba-Wuxi-Badong respectively. Within the present basin, Central Sichuan palaeouplift, and depressions of Langzhong-Tongjiang, Jiangjin-Fengjie constitute of the palaeogeographic pattern of "one uplift and two depressions", while in the south an inherited depression region named "Luzhou-Junlian-Zhaotong" still existed. To the northeast and southwest corner, Qinling ocean and Jiangnan Basin are connected. Strikingly impacted by the palaeogeographic pattern, a land-tied carbonate platform was found to cover the main body of the Sichuan Basin within the platform-ramp-basin sedimentary system, which may further be divided into mixed tidal flat, restricted to evaporate platform, semi-restricted to restricted platform and platform margin. Furtherly, by analyzing the terrigenous clasts content and its direction, the content distribution pattern of limestone, dolostone and gyprock elucidating the waters localization and the relative locations of the platform margin, and the grain-shoal content, it clarifies the distribution pattern of the lithofacies paleogeography during Longwangmiao period, establishing the sedimentary model successively. The results indicate that, the open ocean side of the flat tidal, underwater palaeouplift and the periphery areas of Luzhou-Junlian-Zhaotong are preferable zones for grain shoal development. The filling and joining during the early stage of early Cambrian results in a unified palaeouplift in Leshan and Ziyang, while the north-south staphrogenic Trough in Tongzi-Junlian during the end of Ediacaran and Early Cambrian turned into the Luzhou-Junlian-Zhaotong depression within the platform. Two depression zones in the south have been verified by drilling holes of Well Dengtan 1 and Well Taihe 1. The results may help to guide and broaden the exploration field of karstified grain-shoal reservoir study.

Key words: grain shoal; palaeouplift; land-tied carbonate platform; platform margin; depositional model; palaeogeographic pattern